High potential for weathering and climate effects of non-vascular vegetation in the Late Ordovician

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It has been hypothesized that predecessors of today’s bryophytes significantly increased global chemical weathering in the Late Ordovician, thus reducing atmospheric CO2 concentration and contributing to climate cooling and an interval of glaciations. Studies that try to quantify the enhancement of weathering by non-vascular vegetation, however, are usually limited to small areas and low numbers of species, which hampers extrapolating to the global scale and to past climatic conditions. Here we present a spatially explicit modelling approach to simulate global weathering by non-vascular vegetation in the Late Ordovician. We estimate a potential global weathering flux of 2.8 (km3 rock) yr−1, defined here as volume of primary minerals affected by chemical transformation. This is around three times larger than today’s global chemical weathering flux. Moreover, we find that simulated weathering is highly sensitive to atmospheric CO2 concentration. This implies a strong negative feedback between weathering by non-vascular vegetation and Ordovician climate.

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lichens and bryophytes have been shown to significantly enhance weathering of the surface rocks on which they grow compared with abiotic conditions. They increase both physical as well as chemical weathering rates: lichen rhizines and bryophyte rhizoids break up the rock surface, followed by dissolution of the minerals through different organic acids, alkanolysis and chelating agents released by the organisms. Measurements of weathering intensity below lichens indicate an enhancement by around one order of magnitude compared with abiotic conditions for cool mountainous regions, such as Norway, as well as for warm arid regions, for example, Lanzarote. In moist tropical regions an enhancement of two orders of magnitude has been observed. These measurements are complemented by studies based on mini-watersheds in the northeastern United States, which quantify directly the amount of chemical elements exported from the rock and which confirm the enhancement effect (factor 2–16). Bryophytes, too, have been proven to enhance weathering significantly under field conditions in Canada as well as in microcosms. At the global scale, however, it is difficult to assess the role that lichens and bryophytes play in weathering based on these local findings. The global distribution of rocks covered by lichens and bryophytes is poorly known. Several studies have concluded that lichens and bryophytes could not significantly enhance the global weathering flux, since vascular plants show a considerably higher enhancement of weathering. Other studies, however, suggest that enhancement of weathering by lichens and bryophytes is still sufficiently large for global effects.

The uncertainty regarding the relative importance of lichens and bryophytes for the global weathering flux limits climate reconstructions of the geological past. On the basis of field observations, it has been concluded that the emergence of large vascular plants with rooting systems in the Devonian led to a large increase in global chemical weathering rates. Including this effect in a global geochemical model causes a decrease in atmospheric CO₂ based on the well-known silicate-weathering feedback.

Predecessors of today’s bryophytes, however, could also have enhanced chemical weathering. Molecular phylogenies support the traditional view that non-vascular (bryophyte) plant lineages (liverworts, hornworts and mosses) all pre-date the origin of vascular plants. The ancestors of modern bryophytes are thought to have evolved during the Mid-Late Ordovician period based on the microfossil record of spores and improvements in molecular clock dating support an early origin. The oldest lichen fossil found so far, exhibiting symbionts and anatomy similar to extant lichens, dates to the Early Devonian; however, lichens could be much older, as shown by a potential symbiosis of fungal hyphae and cyanobacteria from the Late Precambrian. Thus, early non-vascular plants, together with lichens, could also have led to a drawdown in atmospheric CO₂.

It has been hypothesized that the interval of glaciations at the end of the Late Ordovician (peaking in the Hirnantian Stage) could be explained by a decrease in atmospheric CO₂ because of the biotic enhancement of weathering. Therefore, microcosms containing moss growing on mineral surfaces were used to estimate the enhancement of weathering compared with microcosms without moss. The results indicate a considerable increase in weathering for several elements, with highest values for iron (over two orders of magnitude) and phosphorus, with calcium and magnesium weathering amplified by factors of circa 1.5–5 across different silicate rock types. In order to translate the increase in chemical weathering rates into a decrease in atmospheric CO₂, the results from the microcosm experiment were used to calibrate an enhancement factor in a global biogeochemical model. Such an approach neglects the spatiotemporal variation of the global moss cover that results from climate patterns and that clearly controls weathering by moss. Moreover, the enhancement factor is based only on one species of moss, growing in the modern atmosphere, while weathering by lichens and bryophytes has been shown to be species-specific.

In this study, we present a complementary approach, which derives chemical weathering by lichens and bryophytes in the Ordovician from their net primary productivity (NPP), which is calculated by a novel, global process-based non-vascular vegetation model that simulates multiple, physiologically different, artificial species. To translate NPP into a weathering flux it is assumed that lichens and bryophytes dissolve surface rocks to access phosphorus for the build-up of new biomass. The amount of rock that is weathered can then be estimated from the phosphorus equivalent of NPP and the phosphorus concentration of rocks. To account for the limited supply of unweathered rock material in shallow regions, we cap biotic weathering at the erosion rate. Furthermore, we constrain weathering by the transport capacity of runoff for dissolved weathered material. Since these limitations on weathering result in a lower nutrient supply from rocks, we reduce our estimate of NPP accordingly. By representing the species-specific, spatiotemporal dynamics of weathering by lichens and bryophytes in the Late Ordovician, we assess the significance of these organisms for Ordovician weathering and climate.

Throughout the manuscript, we use the term chemical weathering to describe the partial dissolution and chemical transformation of a certain volume of primary minerals into solutes and secondary minerals. This includes all elements, such as calcium, phosphorus, iron and so on, in the primary minerals. We use volume as a basis to compare the rate of chemical weathering to the erosion rate.

Results

Weathering by lichens and bryophytes in the Ordovician. The global pattern of chemical weathering by Ordovician lichens and bryophytes, their NPP and their surface coverage as well as the pattern of the limiting factors on weathering are shown in Fig. 1. The estimates are obtained from a baseline simulation of the non-vascular vegetation model, run at eight preindustrial atmospheric level (PAL; 1 PAL = 280 p.p.m.) of atmospheric CO₂ and 14% of atmospheric O₂ for 600 years to reach steady state, with an initial number of 300 species (see Methods section for details).

Potential NPP of lichens and bryophytes depends only on climatic factors (Fig. 1a). It shows a latitudinal pattern, with maximum values at the equator and high values in the region between 30° and 60°. Between the equator and 30° N and 30° S, respectively, NPP is low and in some regions none of the simulated lichen and bryophyte species are able to survive. This distribution can be explained by water limitation of NPP in (sub)tropical regions and limitation by low temperature and light availability in polar regions (see Methods section). The global pattern of weathering by lichens and bryophytes based solely on their NPP is identical to that of potential NPP, since the conversion factors used to translate NPP into a weathering flux are assumed to be globally uniform (Fig. 1b).

Chemical weathering constrained by both the erosion rate and the transport capacity of runoff for dissolved weathered material is strongest in regions of pronounced relief and high rainfall (Fig. 1c). In areas of flat topography, weathering is limited by erosion, which is directly related to the mean surface elevation above the sea level. In arid regions, weathering is limited by low runoff. The global pattern of weathering thus results from the minimum of the limiting factors NPP, erosion and runoff...
In areas where NPP-based weathering exceeds erosion- and runoff-limited weathering, the supply of phosphorus from rocks is not sufficient to sustain potential NPP, which is consequently reduced (Fig. 1f). The strongest reduction of potential NPP occurs at the equator, because of low erosion rates, and in dry regions at 30°S and at high latitudes, because of low rainfall. At the global scale, NPP-based weathering of 3.6 (km³ rock) yr⁻¹ is reduced by 22% to a realized weathering of 2.8 (km³ rock) yr⁻¹. This is accompanied by a reduction in NPP because of phosphorus limitation in shallow or arid regions. Potential NPP of 18.7 (Gt C) yr⁻¹ is reduced by 23% to 14.4 (Gt C) yr⁻¹ of realized NPP.

Global annual values of lichen and bryophyte weathering, NPP, gross primary productivity (GPP), surface coverage and biomass are listed in Table 1 together with the corresponding values for today’s climate in the presence of vascular plants. Today’s values were calculated by forcing the lichen and bryophyte model with data fields for today’s climate and other current environmental conditions. In addition, shown in Table 1 are observation-based estimates of NPP, GPP, surface coverage and biomass of today’s total terrestrial vegetation and total, that is, biotic plus abiotic, global chemical weathering. Considering only lichens and bryophytes, all values corresponding to the Ordovician are considerably larger than those for
Weathering corresponds to 280 p.p.m. Each dot represents a model simulation at the respective atmospheric CO2 concentration. The black solid line shows weathering enhanced by non-vascular vegetation derived from a 600-year simulation with 300 initial species. The grey dashed line represents abiotic weathering, which includes vascular plants, lichens, and bryophytes. The atmospheric CO2 concentration is 2,240 p.p.m. (8 PAL). The grey areas denote regions where no species are able to survive, notably including the ice sheet mask prescribed from the South Pole to the mid-latitudes.

Table 1 | Global effects of lichens and bryophytes.

<table>
<thead>
<tr>
<th>Simulation</th>
<th>NPP (Gt C) yr⁻¹</th>
<th>GPP (Gt C) yr⁻¹</th>
<th>Cover fraction</th>
<th>Biomass Gt C</th>
<th>Weathering (km³ rock) yr⁻¹</th>
</tr>
</thead>
<tbody>
<tr>
<td>Ordovician</td>
<td>14.4</td>
<td>30.1</td>
<td>0.44</td>
<td>133</td>
<td>2.8</td>
</tr>
<tr>
<td>Today</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Lichens and bryophytes</td>
<td>2.6</td>
<td>5.4</td>
<td>0.14</td>
<td>51</td>
<td>0.044</td>
</tr>
<tr>
<td>Total vegetation</td>
<td>60 (ref. 25)</td>
<td>123 (ref. 68)</td>
<td>0.74 (ref. 69)</td>
<td>442 (ref. 70)</td>
<td>0.85 (ref. 54)</td>
</tr>
</tbody>
</table>

GPP, gross primary productivity; Gt C, gigatons of carbon; NPP, net primary productivity.

Annual global values of NPP, GPP, surface coverage, biomass and weathering for the Ordovician baseline simulation (see Methods section) and for today’s climate. Note that we converted the weathering estimate54 by using an average density of surface rocks22.

Figure 2 | Sensitivity of weathering to environmental factors. (a) Global chemical weathering as a function of atmospheric CO2 concentration. One PAL corresponds to 280 p.p.m. Each dot represents a model simulation at the respective atmospheric CO2 concentration. The black solid line shows weathering enhanced by non-vascular vegetation derived from a 600-year simulation with 300 initial species. The grey dashed line represents abiotic weathering, calculated by a GEOCARB approach. (b) Effect of a continental-sized ice sheet south of 30° S on the global pattern of weathering for the Late Ordovician (450 Ma), derived from NPP of lichens and bryophytes constrained by erosion and runoff. The map shows average values over the last 50 years of a 600-year simulation with 300 initial species. The atmospheric CO2 concentration is 2,240 p.p.m. (8 PAL). The grey areas denote regions where no species are able to survive, notably including the ice sheet mask prescribed from the South Pole to the mid-latitudes.

Sensitivity of weathering to environmental factors. Estimated global chemical weathering by lichens and bryophytes in the Ordovician shows a strong dependence on atmospheric CO2 concentration, as illustrated in (Fig. 2a). The increase in weathering with increasing CO2 results from the positive effect of high atmospheric CO2 concentrations on NPP. While the relation looks exponential at values of CO2 lower than 8 PAL, it seems to change at high values to a logarithmic one. Below we discuss possible reasons for the strong dependence of NPP on atmospheric CO2.

We, furthermore, show weathering for abiotic conditions in (Fig. 2a), which is calculated using the GEOCARB III model14. Abiotic weathering is in general considerably lower than weathering enhanced by non-vascular vegetation. However, abiotic weathering does not show saturation at very high values of CO2, in contrast to biotic weathering.

To account for the existence of a continental-sized ice sheet south of 30° S in the Hirnantian Stage24, an additional vegetation model run is performed with a corresponding glacier mask and updated climatic forcing fields (Fig. 2b). Although the ice sheet does not cover the regions of highest NPP and weathering at the equator and close to 30° S, it covers vast regions of Gondwana and, therefore, reduces weathering considerably from 2.8 to 1.5 km³ yr⁻¹ and NPP from 14.4 to 7.6 Gt yr⁻¹.

Sensitivity of weathering to number of artificial species. Figure 3 shows that global weathering by lichens and bryophytes increases with the number of initial artificial species used in a simulation and remains constant for larger numbers of initial species. This results from an undersampling of the very large space of possible combinations of physiological parameters. Species that are capable of high NPP and, consequently, strong weathering, are probably not generated if the number of initial species is limited.
species is too low. This means that, at the beginning of the colonization of the land surface by non-vascular plants, biotic weathering could have been lower because of very low diversity and productivity.

**Discussion**

In this study we estimate chemical weathering in the Late Ordovician by early forms of lichens and bryophytes. We quantify their NPP during this period with a process-based vegetation model and then translate the NPP into a weathering flux, additionally accounting for limitation of weathering by low erosion and runoff.

Our main finding is a huge weathering potential of Ordovician lichens and bryophytes, more than three times higher than today's combined abiotic and biotic global chemical weathering flux. This value results from a 60-fold increase in weathering by these organisms compared with today's value (Table 1), which could be explained by the following factors: first, lack of vascular plants in the Ordovician allows lichens and bryophytes to cover the entire area available for growth of vegetation. Second, Ordovician lichens and bryophytes have increased access to mineral surfaces because of the absence of competing weathering by vascular plants. Finally, high Ordovician atmospheric CO₂ has a strong increasing effect on NPP of lichens and bryophytes and, consequently, on NPP-based weathering.

In the following, we discuss the plausibility of our estimates, starting with the large value of NPP for the Ordovician, which is \( \sim 25\% \) of today's total terrestrial NPP. The main factor for such a high productivity is the strong CO₂ sensitivity of simulated NPP, combined with high levels of atmospheric CO₂ in the Ordovician. The CO₂ response can be analysed in the model for individual artificial species, by varying ambient CO₂, but otherwise keeping constant boundary conditions, such as water content, temperature and radiation. To find appropriate species, we choose two regions with similar climate, one from a simulation with today's climatic and CO₂ forcing and another from an Ordovician simulation with a high atmospheric CO₂ value of 24 PAL. From these two regions we select the dominant species, shown in Fig. 4a. It can be seen that for the high Ordovician CO₂ level, the model favours a species that has a higher photosynthetic capacity than the species dominant under today's CO₂, given otherwise similar climatic conditions. This seems plausible from the viewpoint of natural selection and leads to a higher NPP in the Ordovician simulation. Most dynamic vegetation models, however, do not implement such a large physiological flexibility, but rely on a few plant functional types with fixed photosynthesis parameters. They cannot adapt to strongly differing environmental conditions. Hence, we attribute the strong sensitivity of NPP to CO₂ to the physiological flexibility of our model.

Experiments confirm that the CO₂ sensitivity of lichens and bryophytes is species-specific (Fig. 4b). The species differ with regard to the slope of the CO₂ response as well as the ambient CO₂ concentration, at which photosynthesis saturates. We want to point out that Fig. 4 does not represent an evaluation of the model, it rather illustrates the large variety in CO₂ responses of different lichen and bryophyte species. This supports the possibility of high NPP of early lichens and bryophytes in the Late Ordovician.

Two of the experimental studies considered here measure NPP of bryophytes at CO₂ concentrations up to 8 PAL (see also Methods section). They find a two to threefold increase in NPP, whereas our artificial species from today shows a sixfold increase,
relative to NPP at 360 p.p.m. of CO$_2$. This may be explained by the fact that the artificial species has a lower photosynthetic capacity than the bryophytes from the experiments, maybe to avoid high respiration costs. As a consequence, NPP of the artificial species increases only slowly with rising ambient CO$_2$ compared with the bryophytes (see Methods section for details). For the same reason, however, NPP of the artificial species saturates at a higher CO$_2$ concentration than the bryophyte, leading to a larger CO$_2$ sensitivity. We do not want to exclude the possibility that our model is too flexible regarding the shape of the CO$_2$ response. We therefore test a minimum CO$_2$-sensitivity scenario, where potential NPP is reduced by 50%. Since weathering depends also on environmental limits in our approach, this translates into a 40% reduced estimate of global chemical weathering for the Ordovician of 1.6 (km$^3$ rock) yr$^{-1}$.

Several studies substantiate the occurrence of high non-vascular NPP in the geological past. On the basis of estimated Ordovician soil CO$_2$ production, it has been concluded that the above-ground productivity fuelling this CO$_2$ production must have been similar to today’s (ref. 26). Hence, non-vascular productivity could have been as high as the vascular one. The course of $\delta^{13}$C through the Phanerozoic shows a marked increase in the Late Ordovician consistent with a new source of organic carbon on land and concomitant increase in global organic carbon burial$^{27}$. Moreover, pre-vascular organic carbon deposits were already large, although they are not in the form of coal$^3$. The occurrence of coal before the Devonian is unlikely, albeit difficult to assess because of the sparse occurrence of terrestrial sediments preserved from before the Devonian$^{16}$.

An alternative estimate of NPP for the Ordovician$^{28}$ has been made by multiplying today’s global cryptogamic productivity on ground$^{12}$ with a value of 1.8, which is the ratio of today’s total vegetated area to the area covered by cryptogams. The estimated value of 4.3 (Gt C) yr$^{-1}$ is markedly lower than our estimate, which can be explained by simplifying assumptions in this study, such as a linear relationship between area and productivity and no CO$_2$ fertilization.

Our estimate of realized NPP for the Ordovician is around six times higher than today’s NPP of lichens and bryophytes. However, simulated chemical weathering by the organisms is enhanced by a factor of 60 in the Late Ordovician compared with today’s value. We suggest that the absence of vascular plants in the Ordovician explains the remaining part of the difference in chemical weathering rates. Vascular plants are likely more efficient in enhancing weathering rates than lichens and bryophytes$^{22-29}$. As a consequence, lichens and bryophytes are usually separated from unweathered rock material by a layer of soil in ecosystems that are dominated by vascular plants, such as forests and grasslands. In these ecosystems, growth of lichens and bryophytes on bare rocks is confined to boulders and cliffs, which cover only a small fraction of the land surface. This strongly limits the contribution of lichens and bryophytes to total chemical weathering. In other ecosystems, such as cold and warm deserts, lichens and bryophytes are often the dominant vegetation. However, low runoff largely limits chemical weathering by the organisms in these regions. To assess the effect of soils on weathering by today’s lichens and bryophytes, we assume that all simulated growth in the model may lead to weathering, not only the part that takes place on bare soil. This increases weathering by today’s lichens and bryophytes by a factor of $\sim$10, which, in combination with increased atmospheric CO$_2$, explains the strong increase in weathering from today to the Late Ordovician (Fig. 5).

Given the large value of today’s total terrestrial NPP compared with the Ordovician, the current global chemical weathering flux appears to be quite small (Table 1). This implies that, for a given amount of biotic weathering, today’s vascular vegetation exhibits a much higher productivity than Ordovician lichens and bryophytes at the global scale. We explain this finding by a high nutrient recycling ratio of modern ecosystems of $\sim$50:1 (ref. 30), which results from the deep rooting systems of vascular plants and associated mycorrhiza. Although lichens and bryophytes are able to resorb phosphorus from senescing tissue, they most likely cannot access phosphorus that has been leached or released from decaying biomass and transported by water to deeper soil layers. In our model, we therefore assume a low recycling ratio of 5:1 for Late Ordovician ecosystems. This value is based on measurements of both weathering rates of lichens on rocks as well as their productivity in desert regions$^{20}$ (see Methods section). Our estimates of chemical weathering and NPP are sensitive to the recycling ratio. In the unlikely case that Ordovician ecosystems had the same high recycling ratio of 50:1 compared with modern ecosystems, weathering decreases to 0.34 (km$^3$ rock) yr$^{-1}$ and NPP increases to 17.6 (Gt C) yr$^{-1}$. For a value of 1:1, meaning no recycling, weathering increases to 3.7 (km$^3$ rock) yr$^{-1}$ because of the organisms’ higher phosphorus requirement, while NPP is reduced to 3.9 (Gt C) yr$^{-1}$ owing to nutrient limitation.

It is likely that vascular plants are more efficient in enhancing weathering than lichens and bryophytes$^{22-29}$. However, our estimate of global chemical weathering in the Late Ordovician

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**Figure 5 | Weathering for today’s climate.** Global pattern of chemical weathering by lichens and bryophytes for today’s climate, constrained by their NPP, erosion and runoff. (a) NPP-based weathering is constrained in each grid cell to the fraction of bare rock surface. (b) NPP-based weathering may take place on the whole land surface area. The map shows average values over the last 50 years of a 600-year simulation with 300 initial species. Runoff is derived from annual average values of climate variables$^{20}$. Note the differing ranges of the colour scales.
is higher than today’s weathering (Table 1). One explanation is that the environmental limits, erosion rate and runoff allowed for a much higher global chemical weathering rate in the Late Ordovician compared with today’s value. Thus, weathering by vascular plants today may be largely limited by environmental factors, and not by the ability of the plants to enhance weathering. In parts of the Amazon basin, for instance, the weathering capacity of the vegetation is much higher than the weathering allowed by the low erosion rate, leading to thick, highly weathered soils, but a relatively low value of chemical weathering31,32. In contrast, chemical weathering in the Ordovician is limited by NPP in many regions (Fig. 1). This means that hypothetical vascular vegetation in the Ordovician could increase weathering to a value of 3.8 (km$^3$ rock) yr$^{-1}$, which is the maximum chemical weathering possible under erosion- and runoff limitation.

When early lichens and bryophytes colonized the land surface in the Ordovician, they could well have used their full potential to enhance chemical weathering, until weathering became limited by erosion in many regions and a certain amount of organic material including phosphorus had accumulated in some ecosystems, facilitating a certain degree of recycling. This rationale supports a scenario of strong weathering by lichens and bryophytes leading to a large, but temporally limited release of phosphorus into the oceans and thereby to a short period of cooling because of enhanced marine productivity11. The short duration of the phosphorus release is explained by an ecosystem shift to more efficient nutrient recycling caused by nutrient limitation.

Both the erosion- and the runoff-based limits to chemical weathering are maximum rates, which means that the true rate of chemical weathering may be lower for some reason, such as a strong kinetic control on the weathering reactions because of low residence time of water in the soil, for instance. Several observations indicate, however, that reaction kinetics are often not the limiting factor for chemical weathering at the global scale33. First, chemical weathering does not show a very clear dependence on temperature at large scale. Second, chemical weathering increases with runoff for low and intermediate values of runoff, which should not be the case if the weathering reaction is far from equilibrium. Finally, chemical weathering rates determined in laboratory experiments exceed those measured in the field by up to two orders of magnitude, which suggests that factors other than reaction kinetics may limit weathering rates in many regions of the world. However, in some areas, such as humid mountainous regions, kinetic limitation of weathering reactions may be important. This is also supported by our alternative estimate of today’s global chemical weathering of 1.3 (km$^3$ rock) yr$^{-1}$ based on the GEM-CO2 model, which is smaller than the 1.8 (km$^3$ rock) yr$^{-1}$ from our limit-based approach. To account for this potential overestimation, we assume that the relations between runoff, lithology and chemical weathering established in GEM-CO2 for today are the same for the Ordovician. We then scale our estimate of Ordovician weathering by the ratio of today’s estimates from GEM-CO2 and the limit-based approach. This would reduce global chemical weathering in the Ordovician by 28% to 2.0 (km$^3$ rock) yr$^{-1}$.

The weaker increase in weathering and NPP at high CO2 values in (Fig. 2a) is because of the saturating Michaelis–Menten kinetics of the Farquhar scheme. Abiotic weathering, in contrast, does not show saturation. It increases almost linearly with atmospheric CO2 because of an exponential temperature dependence, which is largely compensated by the logarithmic dependence of temperature on CO2. The stronger than linear increase in biotic weathering and NPP for low CO2 values is probably due to the retreat of continental ice sheets at rising CO2 and a large increase in global land surface temperature. In addition, abiotic weathering shows a nonlinear response to increases in atmospheric CO2 below values of 8 PAL, since the relation between CO2 and land surface temperature is not logarithmic anymore at low CO2 values. The reason for this is a nonlinear increase in sea-surface temperatures with CO2 because of the complex ocean dynamics in the Fast Ocean Atmosphere Model (FOAM)34, which forces the Laboratoire de Météorologie Dynamique Zoom (LMDZ) atmospheric circulation model and, thereby, land surface temperature. The continental ice sheets result from a positive snow balance in the lichen and bryophyte model, which is calculated as a function of snowfall, surface temperature and a rate of lateral glacier movement. This snow scheme, however, is calibrated for the Holocene. For 8 PAL of CO2, for instance, it estimates only a small ice sheet near the south pole (Fig. 1), while there are indications that the ice sheet was much larger35. Given the large uncertainty concerning the extent of the Ordovician ice sheet, we think that using the default snow scheme in the model is appropriate.

Our results indicate that weathering by lichens and bryophytes is sensitive to the diversity of plants growing on the land surface (Fig. 3), and this raises the possibility that fluctuations in the diversity of lichens and bryophytes through geological time could have major biogeochemical consequences. Owing to a lack of macroscopic plant remains in rocks of Ordovician age, plant diversity at this time can be examined using fossilized spores35. These fossils provide the first widely accepted evidence for land plants during the Dapingian-Darriwilian interval36, and preserve evidence of an adaptive radiation of vascular plants in the Palaeozoic37. In light of our modelling results, shown in Fig. 3, we suggest that the magnitude of the weathering flux could have increased through the Ordovician as a consequence of the evolutionary radiation of primitive land plants. Thereby, our study highlights a possible link between the process of diversification and the biogeochemical processes that regulate Earth’s climate.

Fossil cryptospores are not sufficiently abundant in the fossil record at this time to permit a detailed comparison between simulated vegetation and data on ancient vegetation derived from fossils. Nevertheless, recent work has demonstrated that cryptospores are found on all major Ordovician palaeocontinents and that, consequently, early land plants showed a global distribution before the Silurian18,38. Our simulations produce terrestrial vegetation that has such a global distribution (Fig. 1a), and in this respect our simulations are compatible with the available palaeobotanical data37,38.

In summary, our estimates imply that the Ordovician non-vascular biosphere had a similarly high capacity for biotic weathering than today’s biosphere. That should be taken into account for understanding and modelling palaeoclimatic evolution. The colonization of the land by early forms of lichens and bryophytes in the Mid-Late Ordovician16,39 could have caused a large decrease in atmospheric CO2 concentration at the end of the Ordovician through the enhancement of silicate weathering. The results from this study could be used as a basis for further analyses of the feedback between weathering and climate in the geological past. By using a coupled biogeochemical model, the effect of weathering on atmospheric CO2 concentration could be quantified. Subsequently, the CO2 sensitivity of weathering by lichens and bryophytes derived here could be used to estimate the impact of changing CO2 and climate on the biotic enhancement of weathering.

Methods

The non-vascular vegetation model. The process-based non-vascular vegetation model used in this study is a stand-alone vegetation model that uses gridded climatic fields to compute NPP as the difference between photosynthesis and
respiration. Similar to many dynamical vegetation models, photosynthesis is calculated by the Farquhar scheme and respiration via a Q10 relationship. These approaches are combined with lichen/bryophyte-specific characteristics, for example, a passive control of the water status (poikilohydry) or the decrease in CO₂ diffusivity with increasing water content. In contrast to most global vegetation models, which use plant functional types, the lichen and bryophyte model explicitly represents physiological variation: the ranges of observed values of physiological parameters are sampled by a Monte-Carlo algorithm to generate artificial species. These species do not correspond directly to certain lichen and bryophyte species in the real world, but we assume that they represent the diverse physiological strategies of real lichens and bryophytes. This is a novel approach to include physiological variation in models of terrestrial vegetation, which is also used, for example, in the successful Jena Diversity-Dynamic Global Vegetation Model.

The model is run for all species at each grid point of the climatic fields and the performance of the different species is analysed. Those species that cannot maintain their biomass are assumed to die out at the respective grid point and the remaining ones are used to compute NPP. This is performed by weighting each surviving species by its NPP accumulated over the simulation.

Note that processes related to uptake and loss of nutrients are not explicitly modelled. Instead, it is assumed that the costs of weathering agents and nitrogen fixation are included in the respiration term, which is based on observed values.

**Orдовик и окружающие факторы.** Лichen и бройфон моделируют поверхность и NPP как функцию климатических условий: проникающие лунообразные и длинноволновую радиацию, осадки, снегопад, атмосферную температуру, ветер и относительную влажность. Эти значения рассчитаны с помощью атмосферной модели глобального циркуляции модели LMDZ для позднего Ордовика, 450 миллион лет назад.

Модель LMDZ — модель-предшественник на геологический класс атмосферной модели глобального циркуляции модели, работающей как атмосферная компонента Института Симон Лапласа (IPSL) земной системы модели, которая на основе современных и динамических упрощений. Значения средних моделей в этом исследовании, предоставляя решение размерности 3.75° × 1.875° с 39 вертикальными уровнями (21 в тропосфере) на прямоугольном полигоне. LMDZ поддерживает с поверхности, и температуры поверхности (East Ocean Atmosphere Model (FOAM)), также работает на 450 Ma, чтобы учесть эффекты влияния океана, атмосферы и моря. Палеогеография для этих симуляций соответствует Блэкйдж.

**Configuration**. Further details about the FOAM set-up can be found in a previous study. LMDZ is run for 20 years to reach the atmospheric steady state. The climate variables needed by the lichen and bryophyte model are provided in the form of monthly values averaged over the last 5 years of the LMDZ simulation, see Fig. 6 for examples. A comparison of the areas of simulated low rainfall and high temperature with the spatial distribution of indicators of arid climate shows reasonable agreement.

The lichen and bryophyte model runs on an hourly time step to resolve fast processes such as dew formation, which are important for correctly simulating the organisms’ NPP. Hence, to make the monthly climate variables from LMDZ usable for the lichen and bryophyte model, they are interpolated to hourly variables by a stochastic weather generator that was developed for this study (see Supplementary Fig. 1 and Supplementary Note 1).

In addition to the climatic fields, the lichen and bryophyte model requires information about the fraction of surface area that is available for growth, the frequency of disturbance events, such as fires, and the roughness length of the surface. The latter is needed to compute fluxes of heat and water between the lichen or bryophyte surface and the atmosphere. For the Orдовик, we assume that the whole land surface area is available for growth of lichens and bryophytes, except for regions covered by glaciers. We know little about the disturbance regime in the Orдовик, other than that there is no evidence of fires in the form of fossil charcoal, suggesting that oxygen was less than 17% of the atmosphere. Hence, we choose a relatively large disturbance interval of 100 years, which is a typical value for the boreal forest biome. The surface roughness length is set to a value of 0.01 m, corresponding to a desert or tundra ecosystem.

**Model setup.** As baseline simulation for this study, the lichen and bryophyte model is run with the interpolated hourly climatic fields from LMDZ and an atmospheric CO₂ concentration of 8 PAM, which equals 2,240 p.p.m. This CO₂ level is close to the threshold required to trigger Late Ordovician glaciations in several climate models. On the basis of the absence of fossil charcoal in the Ordovician and global geochemical modelling studies, atmospheric O₂ is set to 14% in our simulations. The initial number of artificial species is set to 300, since this amount of species sufficiently represents physiological variation (see Fig. 3). The model is run for 600 years to achieve a state where the number of surviving strategies remains constant for more than 100 years. We assume that this is a sufficient condition for a steady state. Since only 1 year of interpolated LMDZ output is available.
available, we repeat this year 600 times. The estimates of the study are based on the averaged output of the last 50 years of the simulation.

Derivation of chemical weathering from NPP. To derive chemical weathering by lichens and bryophytes from their NPP, we use a method from a previous study51. First, the phosphorus requirement associated with a given NPP is calculated from the C:P ratio of lichen and bryophyte biomass. This phosphorus requirement is then translated into a potential uptake of phosphorus from surface rocks, thereby accounting for resorption of phosphorus from decaying biomass46 and loss of phosphorus through leaching47. The values of this ratio, resorption efficiencies, and losses to leaching vary considerably between different species of lichens and bryophytes. Here we use medium values of these properties22, and we assume that today’s lichens and bryophytes do not differ much from their Ordovician predecessors in this respect. Finally, the amount of rock weathered by lichens and bryophytes is estimated from their uptake of phosphorus and assumes a phosphorus concentration in the Ordovician of 0.9% as the limit-based approach. We apply this way estimates of both biotic and abiotic weathering are based on the same CO2 concentration estimate. The dust could therefore represent a source of phosphorus in arid regions, where no runoff is generated and dewfall is the only source of water for lichens and bryophytes.

Evaluation of erosion- and runoff-limited chemical weathering. We use our limit-based approach to estimate today’s global chemical weathering flux. To be consistent with our estimate for the Ordovician, we use the same globally uniform mean values of soil mineral composition and soil CO2 concentration. We determine runoff for each grid cell from the Budyko curve, where we calculate potential evaporation by the equilibrium evaporation approach53. Annual mean values of radiation, temperature, and precipitation are computed from today’s climatic fields20. The simulated potential global calcium flux amounts to 10.7·1012 mol yr\(^{-1}\), which compares well with the 9.3·1012 mol yr\(^{-1}\) estimated by a more complex model that represents the spatial distribution of lithological classes and includes a process-based scheme for soil CO249. Subsequently, we take the maximum of the runoff-based calcium flux and the flux supported by erosion, where we convert the rock material flux associated with erosion to a potential calcium flux using average mineral properties. This reduces the estimated calcium flux to 7.7·1012 mol yr\(^{-1}\), which corresponds to a global chemical weathering of 1.8 (km3 rock) yr\(^{-1}\). This value is around twice the value based on measurements in the world’s rivers54, converted to rock volume using an average rock density of around 2,500 kg m\(^{-3}\). We explain this by two reasons: First, the measurements likely underestimate the rock volume affected by chemical weathering, since some rock types form insoluble clay minerals during weathering, which are not captured by the measurement method (total dissolved solids). Second, kinetic limitation of weathering reactions may be important in some regions, such as mountain areas in the humid tropics, for instance, leading to an overestimation of chemical weathering by the limit-based approach.

To test these potential explanations we apply the GEM-CO2 model21 that predicts chemical and abiotic weathering as a function of runoff for different lithology classes. These classes are the same as those used in our approach. On the basis of average mineral properties we estimate a global calcium flux of 5.5·1012 mol yr\(^{-1}\), which translates into a rock volume of 1.3 (km3 rock) yr\(^{-1}\). This suggests that clay formation and kinetic limitation contribute roughly equally to the overestimation of chemical weathering by the limit-based approach compared with observations from the world’s rivers. Keeping these uncertainties in mind we believe that the erosion- and runoff-based limits on chemical weathering represent a reasonable first approximation to the actual rate of global chemical weathering and should be suitable to constrain our estimate of chemical weathering in the Late Ordovician to realistic values.

To distinguish between biotic and abiotic contributions to the global chemical weathering flux, we quantify Ordovician abiotic weathering using the GEOCARB III model34. This means that we scale today’s global calcium flux to the Ordovician calcium flux, thereby assuming abiotic conditions in the Ordovician. As a first step, we scale the today’s calcium flux based with chloride concentrations as well as carbonates with a spatially explicit version of the GEM-CO2 model, since GEOCARB uses different equations for the two fluxes. We estimate 2.1·1012 mol yr\(^{-1}\) of carbonate weathering and 3.4·1012 mol yr\(^{-1}\) of silicate weathering, which matches well with the value of 5.5·1012 mol yr\(^{-1}\) for the total calcium flux based on average lithology and mineral properties stated above.

The scaling relationship in GEOCARB III consists of several factors, which describe how changes in temperature, runoff and atmospheric CO2 between today and the Ordovician influence rates of weathering reactions. We use the standard values from GEOCARB III parameterize these factors and we adopt the derivation of runoff from temperature. We do not, however, use the GEOCARB approach to derive Ordovician surface temperature from atmospheric CO2, but instead we use the average land surface temperature estimated by the LMDZ atmospheric circulation model, which is also used to force the non-vascular vegetation model. In this way estimates of both biotic and abiotic weathering are based on the same CO2 concentration. Further, the GEOCARB approach to carbonate weathering further unconstrained. GEOCARB does not distinguish between non-vascular vegetation and abiotic conditions. Hence, the factor of 0.25 integrates studies that estimate weathering below vascular plants to weathering in bare soil as well as studies that compare vascular plants to lichens and bryophytes. For bare soil the factor ranges from 0.25 to around 0.7, but may be higher in areas with weak or no vegetation. Since we do not want to underestimate abiotic weathering, we adopt the GEOCARB III standard value of 0.25 for the Ordovician. Finally, we convert the calcium fluxes to rock volume using average rock properties of carbonate and silicate minerals.

We use the GEOCARB approach for estimating abiotic weathering since our erosion- and runoff-limited approach does not contain an explicit description of the influence of biota on weathering. While the original, more complex version of the limit-based approach includes the impact of root respiration on soil CO2, the atmospheric dust consisted mostly of unweathered rock particles because of the lack of organisms other than lichens and bryophytes, which were capable of transforming significant amounts of rock phosphorus into organic phosphorus. The dust could therefore represent a source of phosphorus in arid regions, where no runoff is generated and dewfall is the only source of water for lichens and bryophytes.

Limits on chemical weathering. In areas where the land surface has a shallow relief and erosion is consequently low, weathering by lichens and bryophytes may become limited by a layer of highly weathered and phosphorus-depleted material blocking the access to unweathered rock substrate. This means that the rate of chemical weathering cannot exceed the rate of exposure of unweathered rock material at the surface, which is equal in the steady state to the erosion rate. We quantify the global distribution of erosion rates as a function of surface elevation above the sea level. We use a method10,22 that relates erosion to the mean relief, and we set the mean relief proportional to the mean surface elevation22. Elevation data for the Ordovician are taken from a global palaeoclimate modelling study34. Our method to estimate chemical weathering implies that a corresponding amount of physical weathering takes place to provide sufficient surface area for chemical weathering. We represent this potential limitation of chemical weathering implicitly by calculating the erosion rate. Only material that has been subjected to physical weathering to some extent can be transported by erosion. Hence, the generation rate of fresh rock material by physical weathering, which equals the erosion rate in steady state, sets the upper limit for the rate of chemical weathering in our approach.

Weathering reactions can only occur at mineral surfaces if water that is saturated with the product of weathering is exchanged by precipitation and runoff. Thus, chemical weathering is limited in steady state to the rate of export of dissolved weathering products. We estimate this export of dissolved material by calculating the calcium flux from the soil as a product of runoff and soil calcium concentration, and then divide by the calcium concentration of rocks to obtain the flux of calcium to rocks. To determine the extent to which chemical weathering is limited by chemical weathering equilibrium approach53. The maximum concentration of calcium in soil water is approximated by the equilibrium with respect to calcite. This means that higher concentrations of calcium in the soil solution are prevented by precipitation of calcite. The calcite equilibrium depends on the partial pressure of CO2 in the soil and on the calcium concentration in the soil solution. The last soil influences the calcium concentration through the relative activity of minerals with and without calcium and the effects of cations other than Ca\(^{2+}\) on the charge balance with HCO3\(^{-}\). Owing to the uncertain spatial distribution of Ordovician surface rocks, we use a globally uniform mineral composition, based on the area-weighted mean for today’s Ordovician. In addition, our rock calcium concentration is based on this mean mineral composition, assuming a calcium content of 40% for calcite and 14% for calcium-feldspar. Since we lack information about soil CO2, we approximate the CO2 concentration based on average lithology and mineral properties stated above. On the basis of today’s atmospheric CO2 concentration, the upper limit for the rate of chemical weathering is based on the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician. The change in land uplift, we use the GEOCARB III value for the late Ordovician.
authors point out that many other factors that are important for enhancing the weathering rate are not considered\textsuperscript{51}. Among these are release of organic acids, retention of water at the rock surface, break up of rock surface by rhizines and rhizoids and supply of carbon to soil microorganisms that perform weathering. Instead of representing the factors explicitly, the limit-based approach rests on the assumption that the factors reduce the kinetic limitation of the weathering reactions to a large extent. As a consequence, chemical weathering is rather limited by climatic and topographic constraints. The fact that the limit-based approach provides a reasonable approximation of realized chemical weathering suggests that kinetic limitation of weathering is indeed not a decisive factor in large parts of the world, such as flat or dry areas. However, without biota, the kinetic limitation of the weathering reactions probably becomes much more important. For this reason, we think it is more appropriate to use a kinetic approach for abiotic weathering, such as the GEOCARB III model.

**Phosphorus limitation of NPP.** In some regions the phosphorus requirement of lichens and bryophytes based on their potential NPP may not be met by the amount of phosphorus available from the substrate, since biotic weathering is limited by low erosion or runoff. In these areas, NPP is phosphorus-limited and we consequently reduce our estimate of NPP by converting the flux of available phosphorus into realized NPP via the C:P ratio of biomass. However, the availability of phosphorus from the substrate may be higher than the amount released by weathering because of additional release of phosphorus by microorganisms that decompose organic matter, or phosphorus remaining in the substrate after leaching. If the water residence time of the substrate is sufficiently long, this additional phosphorus may be taken up by the organisms, resulting in recycling of phosphorus in addition to the effects of resorption and leaching described above, which almost compensate each other. The availability of phosphorus in the substrate probably depends on spatially varying hydrological conditions: rainfall is the main driver for leaching of phosphorus from the organisms, while runoff is necessary to export phosphorus from the substrate. Therefore, phosphorus availability should be lower in humid regions. However, water is necessary for the decomposition of organic matter, which releases phosphorus, leading to lower phosphorus availability in arid regions. In today’s ecosystems, these effects are probably strongly modulated by the effect of roots of vascular plants on soil water-residence time. For simplicity, we assume here a global uniform value of the recycling of phosphorus in the substrate.

The recycling ratio for phosphorus can be estimated by dividing the phosphorus uptake associated with NPP by the phosphorus release from weathering. Since we want to estimate the recycling ratio for the Late Ordovician, we consider those of today’s ecosystems, which are dominated by lichens and bryophytes, such as deserts. Typical values for NPP of lichens and bryophytes in deserts are ~5 (g C) m\(^{-2}\) yr\(^{-1}\), although they show large spatial variation\textsuperscript{22}. Weathering by lichens and bryophytes on rock surfaces\textsuperscript{2} ranges from 0.0004 to 0.01 mm yr\(^{-1}\). Assuming a typical value of 0.002 mm yr\(^{-1}\), the mean values for C:P ratio and rock phosphorus concentration, we calculate a recycling ratio of ~5:1.

We run the lichen and bryophyte model with and without nutrient limitation for today’s climate. We assume a recycling ratio of 5:1 for bare soil, but for regions that are dominated by vascular vegetation we use a recycling ratio of 50:1 (ref. 30). This high recycling ratio lowers the phosphorus requirement of lichens and bryophytes below the amount provided by erosion- or runoff-limited chemical weathering; the reduction of potential NPP thus mostly occurs in desert regions. At the global scale, NPP is reduced by 16% from 3.1 to 2.6 (Gt C) yr\(^{-1}\). In dry areas, the reduction of NPP reduces the average value for the desert biome from 13.7 to 6.2 (g C) m\(^{-2}\) yr\(^{-1}\), which improves the model fit to the typical observed value of 5 (g C) m\(^{-2}\) yr\(^{-1}\) estimated in a previous study\textsuperscript{20}. Note that simulated chemical weathering rates of today’s lichens and bryophytes in desert regions (Fig. 5a) compare well with observed rates on bare rocks\textsuperscript{20}.

**Sensitivity analyses.** We perform a sensitivity analysis to assess the relevance of several uncertainties, which are associated with our modelling approach. To account for the considerable uncertainty regarding the value of Ordovician atmospheric CO\(_2\) content\textsuperscript{3}, we force the lichen and bryophyte model with additional LMDZ climatic fields corresponding to levels of atmospheric CO\(_2\) of 3, 4, 6, 10, 12, 16 and 24 PAL (see Fig. 2a).

We further analyse the sensitivity of simulated NPP to ambient CO\(_2\) concentrations by running the model for individual artificial species for a range of CO\(_2\) concentrations assuming constant light, temperature and water content. We test one species well adapted to today’s climate and another one well adapted to Ordovician climate with an atmospheric CO\(_2\) concentration of 24 PAL (Fig. 4a). To find well-adapted species, we select two regions with similar climatic conditions, one for the Ordovician at 24 PAL and one for today’s climate, and use the parameters of the dominant species in these regions. NPP is calculated in the model as the difference of GPP and respiration. GPP is simulated by the well-established Farquhar scheme, which quantifies photosynthesis as a function of light, temperature and CO\(_2\) concentration within the chloroplasts. The latter depends on diffusion of CO\(_2\) from the atmosphere into the thallus, which is controlled in the model by the limiting effect of water content. The Farquhar scheme depends on several physiological parameters, such as the carboxylation rate of Rubisco, for instance, which vary considerably between the
Table 2 | Sensitivity to model parameters.

<table>
<thead>
<tr>
<th>Scenario</th>
<th>NPP ((Gt C) yr$^{-1}$)</th>
<th>GPP ((Gt C) yr$^{-1}$)</th>
<th>Cover fraction</th>
<th>Biomass (Gt C)</th>
<th>Weathering ((km$^3$ rock) yr$^{-1}$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\Phi_{\text{RR}}$ low</td>
<td>10</td>
<td>23</td>
<td>0.4</td>
<td>102</td>
<td>2.3</td>
</tr>
<tr>
<td>$D_{\text{CO}_2, \text{sat}}$ low</td>
<td>12</td>
<td>25</td>
<td>0.4</td>
<td>124</td>
<td>2.5</td>
</tr>
<tr>
<td>$\eta_{\text{CCM}}$ Low</td>
<td>13</td>
<td>27</td>
<td>0.4</td>
<td>126</td>
<td>2.6</td>
</tr>
<tr>
<td>$t_D$ 250 yr</td>
<td>14</td>
<td>31</td>
<td>0.5</td>
<td>170</td>
<td>2.8</td>
</tr>
<tr>
<td>Baseline</td>
<td>14</td>
<td>30</td>
<td>0.4</td>
<td>133</td>
<td>2.8</td>
</tr>
<tr>
<td>$\eta_{\text{CCM}}$ high</td>
<td>15</td>
<td>31</td>
<td>0.4</td>
<td>135</td>
<td>2.9</td>
</tr>
<tr>
<td>$D_{\text{CO}_2, \text{sat}}$ High</td>
<td>15</td>
<td>31</td>
<td>0.4</td>
<td>134</td>
<td>2.9</td>
</tr>
<tr>
<td>$t_D$ 10 yr</td>
<td>15</td>
<td>27</td>
<td>0.2</td>
<td>40</td>
<td>3.0</td>
</tr>
<tr>
<td>$\Phi_{\text{RR}}$ High</td>
<td>19</td>
<td>38</td>
<td>0.4</td>
<td>153</td>
<td>3.4</td>
</tr>
</tbody>
</table>

GPP: gross primary productivity; Gt C, gigatons of carbon; NPP, net primary productivity.

Global annual values of NPP, GPP, surface coverage, biomass and weathering for the Ordovician baseline simulation (see Methods section) and for upper and lower bounds of four key model parameters. See text for a description of the parameters and the associated upper and lower bounds.
Data availability. The findings in this study are based on climatic fields simulated by the atmospheric general circulation model LMDZ41. Requests for the climate model output can be sent to A.P. (alexandre.pohl@lce.ipsl.fr) or V.D. (yannick.donnadieu@lce.ipsl.fr).

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Author contributions

P.P., A.K. and T.M.L. designed the study. P.P. did the modelling analyses with help and input from A.P. and Y.D.; P.P. wrote the paper with input from T.M.L., A.K., B.W., L.M., C.B. and U.P. in form of literature research and pieces of manuscript text.

Additional information

Supplementary Information accompanies this paper at http://www.nature.com/naturecommunications

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