

Open Research Online

The Open University's repository of research publications and other research outputs

Characterization of the noble gases and CRE age of the D'Orbigny angrite

Conference or Workshop Item

How to cite:

Eugster, O.; Busemann, H.; Kurat, G.; Lorenzetti, S. and Varela, M. E. (2002). Characterization of the noble gases and CRE age of the D'Orbigny angrite. In: 65th Annual Meeting of the Meteoritical Society, 21-26 Jul 2002, Los Angeles, California, USA.

For guidance on citations see [FAQs](#).

© [not recorded]

Version: Not Set

Link(s) to article on publisher's website:
<http://www.lpi.usra.edu/meetings/metsoc2002/pdf/5077.pdf>

Copyright and Moral Rights for the articles on this site are retained by the individual authors and/or other copyright owners. For more information on Open Research Online's data [policy](#) on reuse of materials please consult the policies page.

oro.open.ac.uk

CHARACTERIZATION OF THE NOBLE GASES AND CRE AGE OF THE D'ORBIGNY ANGRITE. O. Eugster¹, H. Busemann¹, G. Kurat², S. Lorenzetti¹, and M. E. Varela³, ¹Physikalisches Institut, University of Bern, 3012 Bern, Switzerland (eugster@phim.unibe.ch), ²Naturhistorisches Museum, Postfach 417, A-1014 Vienna, Austria, ³Dept. de Geologia, Universidad Nacional de Sur, 8000 Bahia Blanca, Argentina.

Introduction: The D'Orbigny angrite, a 16.55 kg stone, was found 1979 in Argentina [1]. Mineralogy and chemistry of this meteorite were characterized in detail [2-6]. A Pb-U-Th age of 4.559 Ga was obtained for pyroxenes by Jagoutz et al. [7]. Here we report results on the noble gas isotopic composition and, in particular, on the cosmic-ray exposure (CRE) age of D'Orbigny.

Cosmogenic component: The nuclides ³He, ²¹Ne, ³⁸Ar, ⁸³Kr, and ¹²⁶Xe are mainly of cosmogenic origin. CRE ages were calculated from the stable isotopes mentioned above, using production rates from Eugster and Michel [8]; the shielding was corrected as for eucrites. The ⁸¹Kr-Kr age was derived using a mean life $\tau_{81}=0.33$ Ma [9] and $P_{81}/P_{83} = 0.44[(^{80}\text{Kr}/^{83}\text{Kr})_c + (^{82}\text{Kr}/^{83}\text{Kr})_c]$ [10]. The results are given in Table 1. The average CRE age is 12.3 ± 0.9 Ma. It differs from that of the other five angrites: Angra dos Reis (55.5 ± 1.2 Ma [11]), LEW 86010 (17.6 ± 1.0 Ma [12]), LEW 87051 (≥ 0.2 Ma [12]), Asuka 881371 (5.4 ± 0.7 Ma [13]), and Sahara 99555 (6.6 ± 0.8 Ma [14]). These differing ejection ages suggest that at least five different sites of the parent body were sampled.

Table 1. CRE ages (Ma) of the D'Orbigny angrite. The average age is 12.3 ± 0.9 Ma.

	T ₃	T ₂₁	T ₃₈	T ₈₃	T ₁₂₆	T ₈₁
Bulk 1 (23.24 mg)	12.0	13.7	11.8			
Bulk 2 (27.04 mg)	11.4	13.8	11.0			
Bulk 3 (379 mg)				10.4	12.5	14.3

Experimental errors (2 σ) of individual ages are ~ 20 %.

Fissionogenic component: We obtain a concentration of $^{136}\text{Xe}_f = 11.4 \times 10^{-12} \text{ cm}^3 \text{ STP/g}$ (> 80 % of total ¹³⁶Xe is fissionogenic). The ratio ($^{134}\text{Xe}/^{136}\text{Xe}$)_f is 0.928 that is close to that for ²⁴⁴Pu fission (0.939 [15]), but differs from that for spontaneous fission of ²³⁸U (0.818 [16]). Our bulk sample contains 0.08 ppm U. Adopting a Xe retention age of 4.56 Ga we calculate $0.3 \times 10^{-12} \text{ cm}^3 \text{ STP/g}$ ¹³⁶Xe_f from ²³⁸U. The high concentration of Xe from short-lived ²⁴⁴Pu is typical for angrites [11, 12] and reflects a relatively short interval between the end of nucleosynthesis and formation of the D'Orbigny material.

Trapped component: D'Orbigny contains no solar noble gases, that is, its material was never exposed on the surface of its parent asteroid. The trapped Kr and Xe is dominated by the cosmogenic and fissionogenic components. Therefore, its type can not be determined.

Acknowledgements: This work was supported by the Swiss National Science Foundation.

References: [1] Meteoritical Bulletin (2001) Meteorit. Planet. Sci. 36, A 296. [2] Kurat et al. (2001) LPS XXXII, abstract # 1737. [3] Varela et al. (2001) LPS XXXII, abstract # 1803. [4] Varela et al. (2001) Meteorit. Planet. Sci. 36, A 211. [5] Kurat et al. (2001) Meteorit. Planet. Sci. 36, A 108. [6] Mittlefehldt et al. (2002) Meteorit. Planet. Sci. 37, 345-370. [7] Jagoutz et al. (2002) LPS XXXIII, abstract # 1043. [8] Eugster and Michel (1995) Geochim. Cosmochim. Acta 59, 177-199. [9] Baglin (1993) Nuclear Data Sheets 69, 267-373. [10] Factor 0.44 from model calculations by B. Lavielle (pers. comm., 2001). [11] Lugmair and Marti (1977) Earth Planet. Sci. Lett. 35, 273-284. [12] Eugster et al. (1991) Geochim. Cosmochim. Acta 55, 2957-2964. [13] Weigel et al. (1997) Geochim. Cosmochim. Acta 61, 239-248. [14] Eugster and Lorenzetti (2001) Meteorit. Planet. Sci. 36, A 55 (revised age). [15] Lewis (1975) Geochim. Cosmochim. Acta 39, 417-432. [16] Eikenberg et al. (1993) Geochim. Cosmochim. Acta 57, 1053-1069.