A short-duration pulse of ductile normal shear on the outer South Tibetan detachment in Bhutan: Alternating channel flow and critical taper mechanics of the eastern Himalaya

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[1] In easternmost Bhutan the South Tibetan detachment (STD) is a ductile shear zone that juxtaposes the Radi (or Sakteng) klippe of the Tethyan Sedimentary Series from underlying high-grade Greater Himalayan rocks. In situ LA-ICPMS U-Th-Pb analysis of metamorphic monazite from the immediate footwall and hanging wall of the STD within the shear zone at the base of the klippe, constrains north vergent normal shear to between 25 and 20 Ma. Coeval thrusting on the Main Central Thrust during this time supports a phase of channel flow–viscous wedge model activity, lasting only ~3 Ma. Geochronologic data from the eastern Himalaya indicate alternating mechanisms for extrusion of the metamorphic core of the orogen from the Late Oligocene through to the Late Miocene, switching from channel flow–viscous wedge behavior to critical taper–frictional wedge behavior, each phase lasting approximately only 2 to 5 Ma. The tectonic evolution of the eastern Himalaya is comparable to central and western Himalayan tectonics during the Early Miocene, but during the Middle Miocene metamorphism and magmatism in the eastern Himalaya migrated toward the orogenic hinterland, a process not widely documented elsewhere in the Himalaya, thus highlighting the need for an orogenic model in three spatial dimensions. Citation: Chambers, J., R. Parrish, T. Argles, N. Harris, and M. Horstwood (2011), A short-duration pulse of ductile normal shear on the outer South Tibetan detachment in Bhutan: Alternating channel flow and critical taper mechanics of the eastern Himalaya, Tectonics, 30, TC2005, doi:10.1029/2010TC002784.

1. Introduction

[2] The Himalayan orogen, formed in response to the collision between India and Eurasia by ~50 Ma (see review in the work of Najman et al. [2010]), represents the archetypal modern continent-continent collision zone. The metamorphic core, the Greater Himalayan Sequence (GHS), is bound by the Main Central Thrust [e.g., Heim and Gansser, 1939] below and the South Tibetan Detachment (STD) [Burchfiel et al., 1992] above. These major crustal-scale fault systems can be traced near continuously along the 2200 km length of the orogen, yet we do not fully understand the evolution of these structures in time and/or space, compromising tectonic models for the Himalaya and for modern orogenesis in general. For example, the evidence for sustained coeval thrusting on the MCT and normal shear on the STD throughout the Early Miocene (23–16 Ma) [Godin et al., 2006, Table 1] supports models of wedge or channel-like extrusion of the GHS [e.g., Jamieson et al., 2004]. Leloup et al. [2010] establish that the switch from movement on the STD in southern Tibet, north of eastern Nepal, to exhumation on north-south trending normal faults at circa 12 Ma is coeval with changes in the basal thrust configuration that is more consistent with wedge extrusion than channel flow, at least by the mid-Miocene. However, the MCT-STD branch line geometry and alternating kinematics recorded on the STD generate contrasting tectonic wedge models, e.g., for the western Himalaya [Yin, 2006; Webb et al., 2007]. Alternatively, a critical taper model based on data from Nepal in the central Himalaya [Kohn, 2008] suggests that the STD played a relatively minor role in the thermomechanics of the orogenic wedge compared to the MCT. It is therefore evident that the timing, duration and kinematics of slip on these faults forms a critical test of the aforementioned models for Himalayan orogenesis.

[3] The tectonic architecture of the Bhutan Himalaya features discrete inner and outer (relative to the hinterland) strands of the STD (Figure 1 [Kellett et al., 2009]), which distinguish the detachment structures from more western Himalayan regions based on the present-day outcrop pattern where typically only one (inner) STD strand is observed. Current data regarding the timing of ductile shear on the STD in Bhutan is limited to the following areas (Figure 1): southern Lingshi klippe, southwestern Ura klippe, Masang Kang [Kellett et al., 2009] and Khula Kangri [Edwards and Harrison, 1997]. Zircon U-Pb data from synkinematic leucogranites indicate that north vergent extensional ductile deformation associated with the outer STD lasted between at least 24 and 16 Ma, and on the inner STD between 15.5 and 11.0 Ma [Edwards and Harrison, 1997; Kellett et al., 2009; Kellett et al., 2010]. However, crystallization ages from synkinematic intrusions can only provide minimum age
Figure 1

2 of 12
2. The South Tibetan Detachment

[5] The STD includes a ductile shear zone usually overlapped at higher structural levels by a discrete normal brittle fault associated with north vergent extensional movement into the upper crust [Burchfiel et al., 1992]. The shear zone, although largely recognized as a north vergent extensional structure, does not record an early (and/or alternating with top-to-the-northeast) top-to-the-southwest kinematic history [Yin, 2006; Webb et al., 2007]. Amphibolite-granulite facies (generally migmatitic and mylonitic) GHS rocks lie in the footwall, juxtaposed against the Tethyan Sedimentary Series (TSS) in the hanging wall. Basal TSS rocks, where preserved between the base of the STD shear zone and the structurally higher brittle STD fault, are metamorphosed to amphibolite facies, e.g., Haimanta Group [Chambers et al., 2009] and Everest Series [Jessup et al., 2008]. Above the brittle STD, TSS rocks are weakly or unmetamorphosed, but do preserve multiple episodes of Eocene-Oligocene deformation in response to initial collision between India and Eurasia [e.g., Godin, 2003]. The age range for ductile shearing on the STD across the entire orogen is 26 to 12 Ma with most data clustering between 23 and 18 Ma [Godin et al., 2006, Table 1]. The majority of these data are granite crystallization ages. Many of these granites intruded during normal shear on the STD, were deformed by ductile shearing, and crosscut by brittle STD faults. These crystallization ages therefore only represent a point in time at which the ductile STD was active, and do not provide constraints on the initiation or cessation of movement on the detachment in the midcrust; data that are key to tectonic modeling.

[6] In the eastern Himalaya, the lower (ductile) and upper (brittle) STD structures described above are recognized in the far north of Bhutan: the contact between the GHS and TSS features discrete steeply north dipping brittle faults that cut mylonitic fabrics associated with previous ductile shear along the STD shear zone [Burchfiel et al., 1992; Edwards et al., 1996]. Yet in more central and southern regions of Bhutan, a high-strain ductile shear zone forms the base of several isolated synformal outliers (klippen) of TSS above gneissic and migmatitic GHS (Figure 1). Following the rationale presented by Kellett et al. [2009], we refer to this high-strain shear zone as the outer STD, and we consider the structures in northern Bhutan to be part of the inner STD system. At the base of the klippen are garnet-staurolite schists of the Chekha Formation. Metamorphic grade and degree of deformation decrease markedly upsection into the TSS where Paleozoic fossils are preserved and sedimentary bedding is the primary planar fabric [Hughes et al., 2011]. Asymmetric shear bands and deformed leucogranite veins in the klippen indicate normal displacement of the hanging wall rocks toward the north [Grujic et al., 2002]. The outer STD as observed at the base of the Ura and Radi klippen lacks evidence for brittle faulting that is observed elsewhere in the Himalaya, including the inner STD in the north of Bhutan.

[7] There are two end-member models for the formation of the inner and outer strands of the STD: (1) critical taper–frictional wedge theory and (2) channel flow–viscous wedge theory, both discussed by Kellett et al. [2009]. In both models the outer STD is abandoned by 12 Ma, and north vergent normal shear is transferred to the inner STD, an active structure from 15.5 to 11 Ma [Edwards and Harrison, 1997; Grujic et al., 2002; Kellett et al., 2009]. In central and eastern Bhutan the termination of the outer STD is considered the result of large-scale folding of the GHS “channel” and subsequent out-of-sequence thrusting between 14 and 10 Ma, thickening the GHS along the northward dipping Kakhtang thrust, and therefore isolating klippen of TSS [Grujic et al., 2002]. This hinterland propagation of exhumation and metamorphism during the Middle Miocene contrasts with the pattern of foreland-directed thrust propagation observed in more western regions of the Himalayan orogen [Hodges, 2000].

3. Geology of the Radi Klippe, Eastern Bhutan

[8] The Radi (or Sakteng) klippe is the easternmost klippe of TSS in Bhutan (Figure 1), and therefore probably also in the Himalayas. The basal Chekha Formation schists feature well-developed, pervasive foliation and crenulation fabrics, formed during lower amphibolite-facies metamorphism. Some garnet porphyroblasts are idioblastic, and these are restricted to quartz-rich crenulation arcs, only weakly wrapped by the matrix foliation and have no internal fabric defined by inclusions (Figure 2f). Other garnet porphyroblasts are subidioblastic and overprinted a crenulated fabric:

Figure 1. (a) Simplified geological map of Bhutan after Bhargava [1995] and Hollister and Grujic [2006]. O-STD and I-STD, outer and inner South Tibetan detachments, respectively. Geological units are as follows: TSS, Tethyan Sedimentary Series; GHS, Greater Himalayan Sequence; LHS, Lesser Himalayan Sequence. Klippen abbreviations are L, Lingshi; TC, Tang Chu; BM, Black Mountain; U, Ura. Plutons are as follows: KK, Khula Kangri; MP, Monlakarchung-Pasalum; GP, Gophu La. (b) Geology of Radi klippe (detail of boxed area in Figure 1a). Metamorphic zones labeled with mineral abbreviations according to Kretz [1983].
Ilmenite inclusions define crenulation limbs (Figure 2e) or possible fluid pathways (Figure 2d). Poikiloblastic staurolite both encloses garnet and overprints the matrix foliation, thus postdating garnet growth and deformation (Figures 2a, 2b, and 2c). Biotite, amphibole and staurolite porphyroblasts have commonly grown across a fine-grained micaceous foliation (Figures 3b and 3c [Gansser, 1983]). Quartzites in the Chekha Formation are not penetratively deformed and preserve cross bedding, providing evidence for strain partitioning (Figure 3a). Further upsection (NE of Radi; Figure 1), metamorphic grade decreases into biotite-grade schists.

Unmetamorphosed, potentially fossiliferous TSS, recognized in the core of other klippen (Figure 1 [Gansser, 1983; Hughes et al., 2011]), are yet to be documented in the Radi klippe. Small (meter-scale) boudinaged leucogranite and pegmatite intrusions cut the main foliation of the Chekha Formation at an oblique angle. At least one kilometer-scale, relatively discordant, granite intrudes higher structural levels in the Chekha Formation (Figure 1b [Bhargava, 1995]).

In contrast to the Chekha Formation schists, the footwall to the outer STD is characterized by mylonitic sillimanite gneiss of the GHS, and the shear zone can be traced based on this lithological contrast (Figures S2, S3, and S4 in Text S1). The GHS rocks in the sampled area (Figure 1b) contain fibrolitic sillimanite in the matrix and matted prismatic crystals on shear surfaces (Figure S8); garnet is also abundant; no kyanite was found, although kyanite is present at structurally deeper levels in the GHS (Figure S6 [Daniel et al., 2003]), and has been found in the uppermost meters of the GHS at the western boundary of the klippe [Grujic et al., 2002]. Tourmaline ± white mica ± garnet leucosome pods and boudinaged veins are common. Within the GHS beneath the western edge of the Radi klippe, subparallel to the base of the klippe (Figure 1b and Figure S5 in Text S1), a 13 Ma leucogranite occurs as a late synkinematic intrusive sheet with 500 Ma augen gneiss xenoliths [Daniel et al., 2003]. Although weakly foliated, this leucogranite appears to have intruded parallel to the strongly layered fabric of the GHS but does not intrude upward into the STD zone.

In the immediate footwall to the outer STD, asymmetric foliation boudinage and shear bands in the GHS gneisses record top down to the northeast. This contrasts with lower levels in the GHS toward the MCT that consistently record south vergent thrusting (e.g., mylonitic augen gneiss with top-to-the-south sense of shear; Figure S7 in Text S1 [Grujic et al., 2002]). At the base of the Chekha Formation, kinematic indicators (stretching lineations, mineral lineations and fibrous minerals) record predominantly north to northeast directed kinematics and minor conjugated south to west directed kinematics, indicating a significant pure shear component during outer STD shearing [Grujic et al., 2002]. Asymmetric pressure shadows on garnet porphyroblasts in the basal Chekha Formation also record northeast vergent normal shear, but at higher structural levels in the klippe kinematic indicators are weakly developed and sense of shear is ambiguous. However,

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1Auxiliary material data sets are available at ftp://ftp.agu.org/apend/tc/2010tc002784. Other auxiliary material files are in the HTML.
asymmetric pressure shadows on garnet porphyroblasts have been observed at one locality which record top-to-the-southwest thrust kinematics.

4. Monazite Analysis

4.1. Sample Descriptions

Samples of the Chekha Formation and the GHS from the outer STD shear zone, Radi klippe (Figure 1b) contain metamorphic monazite grains exclusive to the matrix; that is, they are not present as inclusions in either garnet or staurolite porphyroblasts. In the GHS (sample 67) grains are generally elongate (∼50 × 100 μm), subhedral, and have weak, patchy, or no apparent Y or Th zonation (Figure 4 and Figure S1 in Text S1). In comparison, monazite grains from the Chekha Formation (sample 64) are relatively abundant, euhedral, smaller (∼25 × 40–75 μm) and commonly have distinct Th zonation, corresponding in at least one grain with weak yet discernible Y zonation (Figure 4). For most monazite grains in this study Y is below the detection level. In both samples most of the grains are aligned with the matrix foliation and lineation, indicating synkinematic growth or possibly postkinematic mimetic growth (i.e., growth controlled by preexisting grain arrangements [e.g., Passchier and Trouw, 2005]). However, some grains in the Chekha Formation (e.g., “ChF mon1”; Figure 4) overgrew the matrix foliation and are therefore postkinematic. With regards to other accessory minerals, garnet has inclusions of allanite (both samples) and apatite (Chekha Formation only); apatite is a common matrix phase.

4.2. (U, Th)-Pb Analytical Methods

A UP193SS New Wave Research laser ablation (LA) system was used in conjunction with a Nu Instruments inductively coupled plasma multicollector mass spectrometry (ICP-MC-MS) at the NERC Isotope Geosciences Laboratory, UK, to ablate single spots (15 μm diameter) using laser fluences of 2–3 J cm⁻² (50% power, 5 Hz). Analysis of monazite followed methods similar to those described by Cottle et al. [2007] using the instrumental setup similar to Simonetti et al. [2005]. Inferred ages of monazite growth are based on both U-Pb and Th-Pb measurements. However, instrumental protocols meant that U-Pb and Th-Pb could not be measured simultaneously and were therefore measured sequentially, returning to the exact location to run the remaining Th-Pb analysis sequence. Only one to two analyses per sample grain were possible to ensure sputter cross contamination was avoided and in view of the small grain size of monazite.

Due to the small grain sizes involved, for which it is difficult to accurately measure ²⁰⁸Pb via LA-ICP-MS, we implemented the following common ²⁰⁸Pb correction method to determine radiogenic ²⁰⁶Pb for ²³⁴Th-²⁰⁶Pb ages. In this procedure we (1) assumed that the lower concordia intercepts on the U-Pb Tera-Wasserburg plot represent the correct ²³⁴U-²⁰⁶Pb age of each sample, (2) calculated the
Figure 4. Backscatter electron images of monazite grains from samples 67 (GHS) and 64 (Chekha Formation). Sections were cut perpendicular to foliation and perpendicular to lineation (sample 67) or parallel to the lineation (sample 64). Left insets show enlarged high-contrast BSE images. Right insets show Th ± Y maps (brightness corresponds to concentration). Circles show locations of LA-ICPMS analyses (spot size 15 μm). See auxiliary material for more images of grains from the GHS sample 67 and color versions of the Th and Y maps. Mineral abbreviations according to Kretz [1983].
isotopic composition of this concordant age (i.e., free of common Pb), (3) calculated the fractional contribution of common $^{206}$Pb and $^{207}$Pb that contributed to the measured composition, and (4) used these common Pb contributions, and a model Pb-isotope composition for ~20 Ma [Stacey and Kramers, 1975], to correct the measured $^{208}$Pb for common $^{206}$Pb and calculate radiogenic $^{232}$Th-$^{208}$Pb ages. As estimates of either common $^{206}$Pb or common $^{207}$Pb can be applied in step 3 to calculate common $^{208}$Pb, two age estimates per analysis were generated.

[15] The 554 Ma Manangotry monazite standard [Paquette et al., 1994; Horstwood et al., 2003] and an in-house 54.5 Ma monazite standard, FC-1 were used for Pb/U and Pb/Th normalization, respectively, coupled with a static ablation pattern using identical spot size and fluence ablation characteristics. The FC-1 standard cannot be directly used for U-Pb calibration as it contains variable excess $^{206}$Pb. The overall reproducibility of both standards during the course of these analyses was 2 to 8% (2σ), which has been quadratically propagated into the uncertainties for each spot analysis. These uncertainties are higher than normally observed, partly due to the very small ablation spot size used to obtain the high spatially resolved analyses. Sample data were normalized in sub sessions as appropriate using the standard data over two analytical sessions.

4.3. Results

[15] Isotopic measurements for monazite samples and standards are presented in Table 1 and Data Set S1, respectively. Calculated ages and data plots were generated using Isoplot version 3.40 [Ludwig, 2003]. U-Pb data plotted on a Tera-Wasserburg diagram lie in two linear arrays, one corresponding to each sample: 21.6 ± 0.6 (2σ) Ma, Chekha Formation; 23.4 ± 1.2 (2σ) Ma, GHS (Figure 5). No data were omitted from the regressions on the Tera-Wasserburg plot with the exception of two analyses (5.2, 7) from the GHS sample that plot in between the two groups of data. These data are not included in the regression of data from this sample as they either represent analyses that partially sampled neighboring ilmenite (Figure 4) and were subsequently sample as they either represent analyses that partially sampled these analyses, or similarly, between core and rim analyses, suggesting that zoning was acquired via growth over a very short period of time, i.e., within the errors quoted for each U-Pb age of monazite growth.

[15] Although the interpretation of $^{238}$U-$^{206}$Pb data may be complicated by the presence of excess $^{206}$Pb in Th-rich minerals such as monazite [e.g., Parrish, 1990], we cannot resolve any such effect in our data set. Regardless, any excess $^{206}$Pb is unlikely to account for more than ~0.5 Ma difference in age, since the mean Th-Pb ages (Figure 6 and Table 2) are slightly older but agree within uncertainty, with the U-Pb ages (the opposite would be true if excess $^{206}$Pb was significant). Nevertheless the quoted ages could be slightly too old (i.e., 0.5 ± 0.5 Ma).

5. Discussion

5.1. Monazite Petrogenesis

[19] Since monazite in both our samples is restricted to the matrix, and since allanite is only present as inclusions in garnet, we infer that monazite grew at the expense of matrix allanite following garnet growth [e.g., Janots et al., 2008]. Combined with (1) the lack of evidence for retrograde monazite [e.g., Bollinger and Janots, 2006], (2) the fact that distinct Th zones within one monazite reflect discontinuous reactions during the same metamorphic event [e.g., Foster et al., 2000], and (3) because rimward decrease of Th content in monazite is consistent with subsolidus crystallization during prograde metamorphism [Kohn and Malloy, 2004], all the analyzed monazite grains grew during prograde, near-peak metamorphism. Y zonation in garnet (Figure 7) and monazite in both samples is consistent with this reaction history, as monazite grains are depleted in Y owing to the earlier sequestration of Y in garnet.

[20] The timing of prograde monazite growth is closely associated with the staurolite-in reaction [Corrie and Kohn, 2008], which is consistent with the textural evidence for concurrent, late growth of both monazite and staurolite in the Chekha Formation. Th (+Y)-rich monazite rims in the Chekha Formation, indistinguishable in age from the other analyses of monazite in this sample, may be attributable to either late garnet resorption during initial decompression or the influx of Th (+Y)-rich fluids associated with local leucogranite intrusions.

[21] Assuming sufficient time since collision to allow a lateral thermal gradient to develop between the hinterland and foreland of the orogen [see Kohn, 2008, and references therein], regardless of the orientation of the outer STD from the horizontal, initiation of north vergent shear across the outer STD would have cooled the immediate footwall (GHS) and effectively terminated prograde monazite formation. Therefore the GHS monazite grains are prekinematic or early synkinematic with respect to north vergent STD shear. Consequently, the age of GHS monazite grains constrains the initiation of normal shear on the outer STD to between 25 and 22 Ma (Table 2). Synkinematic to postkinematic monazites from the Chekha Formation, in the immediate hanging wall to the outer STD, date cessation of normal shear on this structure at 22–20 Ma (Table 2). We attribute prograde synkinematic monazite growth in the Chekha Formation to viscous heating, whereas prograde...
posttectonic monazite growth in the same sample is considered the product of advected heat, from the uplifted and therefore hotter GHS below, across the outer STD and into the overlying Chekha Formation. Postkinematic amphibole and staurolite in the Radi klippe (Figures 3b and 3c) are also attributed to advected heat from the GHS. Peak temperature in the Chekha Formation thus postdates peak temperature in the GHS by up to 5 Ma.

### Table 1. (U, Th)-Pb Sample Monazite Isotopic Data

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</table>

### 5.2. Tectonic Evolution

[22] Kellett et al. [2009] outline two models for the tectonic evolution of Bhutan: (1) critical taper–frictional wedge theory and (2) channel flow–viscous wedge theory. Both of these apply to a system in which the STD shear zone is a continuously active structure, since its initiation in the Late Oligocene, through transfer from the outer STD to the inner STD at 15.5 Ma, to cessation in the Middle to Late Miocene.
Figure 5. Tera-Wasserburg plot of U-Pb monazite data from the GHS and Chekha Formation (one analysis of which, included in the regression but with significant common Pb, is therefore not seen at this scale).

Figure 6. Weighted average 208Pb/232Th common Pb corrected monazite ages (Table 1). Error bars are 2σ.

Table 2. Summary of U-Pb and Th-Pb Monazite Ages (±2σ)

<table>
<thead>
<tr>
<th>Isotope System</th>
<th>Chekha Formation</th>
<th>Greater Himalayan Sequence</th>
</tr>
</thead>
<tbody>
<tr>
<td>U-Pb</td>
<td>21.6 ± 0.6 (MSWD = 2.1)</td>
<td>23.4 ± 1.2 (MSWD = 0.48)</td>
</tr>
<tr>
<td>Th-Pb^a</td>
<td>circa 22 Ma</td>
<td>24.5 ± 0.5 (MSWD = 1.4)</td>
</tr>
<tr>
<td>Th-Pb^b</td>
<td>circa 22 Ma</td>
<td>24.5 ± 0.7 (MSWD = 2.5)</td>
</tr>
</tbody>
</table>

^aNote that Chekha Formation Th-Pb ages are approximate, as the data do not define a statistically meaningful single population (Figure 6). Sample 64.
^Sample 67.
^Corrected for common Pb using 206Pb.
^Corrected for common Pb using 207Pb.

shear on the inner STD initiated at 15.5 Ma [Kellett et al., 2009]. The extent to which early (pre-16 Ma) STD shear may also be recorded within the strain history of the upper GHS beneath the inner STD, is as yet undetermined; data from Kellett et al. [2009] do not exclude the possibility that shear across the inner STD may have initiated earlier than 15.5 Ma and thus shared a common movement history with the outer STD. The following discussion considers the implication of a significant hiatus in north vergent normal shear for the geodynamics of the far eastern Himalaya (eastern Bhutan). Although one explanation for this hiatus could be that motion on the inner STD directly north of the Radi klippe predated that on the inner STD, 100–150 km to the west (source region of all the available data), this would imply that STD movement was diachronous (younging eastward by ~5 Ma across Bhutan) for which there is no supporting evidence.

[23] A summary of the geochronologic data from the major shear zones in the eastern Himalaya is presented in
sometime after 15 Ma, and may have been synchronous with
Initiation of south vergent thrusting on the KT occurs
duplex 
MCT overlaps with initiation of the lower Lesser Himalayan
and at 13.4 ± 0.2 Ma. The cessation of shearing across the
normal shear on the inner STD, between 15.5 and 14.5 Ma,
ment across the MCT is also synchronous with ductile

tified between 23 and 21 Ma, and 18 and 16.5. Displace-
coeval shear across the outer STD and MCT can be iden-
MCT, by at least 2 if not 6 Myr (Figure 8). Two periods of
coeval MCT–STD system throughout the Miocene, but the present data set does not unambiguously prove that
this is the case. Note that we refer to Daniel et al. [2003] for
ages relevant to shear on the MCT and KT in eastern
Bhutan, and not to Grujic et al. [2002], who cited inaccurate
age data from Daniel et al. [2003], which at that time was an
unpublished manuscript, in press. Furthermore, Daniel et al.
[2003] present no ages relevant to dating shear on the STD, and the 17–22 Ma maximum age of north vergent shearing
based on a leucogranite crystallization age quoted by Grujic
et al. [2002] is unjustified. We deliberately do not include
the Main Boundary Thrust (MBT) or Main Frontal Thrust
(MFT) in Figure 8, as there are no data to constrain the
timing or amount of slip on these structures in the eastern
Himalaya, so the geodynamics since circa 10 Ma remain
difficult to interpret. Although detrital records from the
western and central Himalaya suggest initiation of the MBT
between 12 and 9 Ma [Meigs et al., 1995; Huyghe et al.,
2001], kinematic reconstructions coupled with detrital
records from the eastern Nepal imply that the MBT was not
activated until <5 Ma [Robinson et al., 2001]. Furthermore
Yin [2006] highlights the fact that the MCT and MBT are
folded in the eastern Himalaya, reflecting a major change in
structural style across the orogen, and therefore age data
from the western Himalaya should not be extrapolated over
1000 km to the eastern Himalaya. The one kinematic
reconstruction for the eastern Himalaya by McQuarrie et al.
[2008] implies only that shortening on the MBT and MFT
occurred after 10 Ma.

Across the eastern Himalaya, north vergent extension
on the outer STD predates south vergent thrusting on the
MCT, by at least 2 if not 6 Myr (Figure 8). Two periods of
coeval shear across the outer STD and MCT can be iden-
tified between 23 and 21 Ma, and 18 and 16.5. Displace-
ment across the MCT is also synchronous with ductile
normal shear on the inner STD, between 15.5 and 14.5 Ma,
and at 13.4 ± 0.2 Ma. The cessation of shearing across the
MCT overlaps with initiation of the lower Lesser Himalayan
duplex [McQuarrie et al., 2008] for a period of 2 to 5 Ma.
Initiation of south vergent thrusting on the KT occurs
sometime after 15 Ma, and may have been synchronous with
ductile normal shear on the inner STD. This episodic history
of thrusting and extension fundamentally supports a critical
taper–frictional wedge theory for the eastern Himalaya,
wherein the taper of the orogenic wedge is continuously
readjusted toward the critical angle through alternate, but at
times coeval, thrust–detachment systems. However, during
periods of coeval MCT–STD (and possibly KT inner STD)
fault movement a channel flow–viscous wedge model may
dominate the thermomechanic behavior of the orogenic
wedge, but ultimately one of the two channel-bounding
faults is deactivated; the MCT at 21 and the outer STD at
~15 Ma (Figure 8).

[25] The first transition in mode, from the critical taper–
frictional wedge 1 to the channel flow–viscous wedge 1 at
23 Ma (Figure 8), may have occurred in response to the
attainment of a threshold in the midcrustal to lower crustal
viscosity in the Himalayan hinterland, such that ductile flow

Figure 7. Major-element (Ca, Mg, and Mn) and Y profiles
across garnet porphyroblasts.

Figure 8. Timeline of movement on major structures in the
eastern Himalaya, divided into west [Edwards and Harrison,
1997; Kellett et al., 2009] and east Bhutan [Stüwe and
Foster, 2001; Daniel et al., 2003; McQuarrie et al., 2008]
(also this study). Bars show known bracketed ages; dashed
bars show poorly constrained ages (e.g., only minimum or
maximum ages of displacement known). O–STD and I–STD,
outer and inner South Tibetan detachments, respectively.
MCT, Main Central Thrust; KT, Kakh t h a n t h r u s t ; l o w e r
Lhd, lower Lesser Himalayan duplex. Black star represents
timing of high-pressure metamorphism (eclogite facies) for
GHS now exposed in NW Bhutan [Grujic et al., 2010].
active, which effectively “turned off” channel flow toward the topographic front, and the system returned to one dominated by a critical taper–frictional wedge (Fig. 8). A change in one or more boundary conditions (such as focal area of denudation, denudation rate, rheology of the deforming crust) at this time would permit such an adjustment of the orogenic wedge.

[26] At 18 Ma the channel flow mode was “turned on” again (channel flow–viscous wedge 2, Figure 8). At this time the lower Lesser Himalayan duplex began to grow [McQuarrie et al., 2008] leading to the formation of a ramp on the MCT. Consequently, the ductile GHS in the MCT hanging wall was destabilized and the GHS channel domed, leading to the termination of motion on the outer ST and initiation of the inner STD [Kellett et al., 2009]. Thus, this channel flow–viscous wedge 2 lasted between 2 and 5 Ma, before the geodynamics changed again at ~15 Ma, either to another critical taper–frictional wedge, or a channel flow–viscous wedge bounded by the inner STD and Kakhtang Thrust. This change coincides with the exhumation of high-pressure GHS rocks at 14–15 Ma, now exposed in the NW of Bhutan [Grujic et al., 2010].

[27] It is apparent that data from elsewhere in the Himalaya, which constrain both initiation and cessation of movement on major structures, are now required for a more detailed comparison between regions of the Himalayan orogen. However, we do not know if the outer ST and klippen of the TSS ever existed in the central or western Himalaya. The TSS klippen (including the Chekha Formation) in the eastern Himalaya may be only locally preserved due to relatively low erosion rates in this region, influenced by the uplift of the Shillong Plateau and an associated decrease in precipitation at circa 5 Ma [Grujic et al., 2006; Biswas et al., 2007], despite relatively high modern-day annual precipitation rates, at least in southern Bhutan, compared to central and western Himalayan regions [Bookhagen, 2010].

6. Conclusions

[28] Monazite U-Th-Pb data, from the outer ST at the base of the Radi klippe in the eastern Himalaya, bracket the initiation (25–22 Ma) and cessation (22–20 Ma) of north vergent normal displacement across the shear zone. South vergent thrusting on the MCT during this time supports a channel flow–viscous wedge model, but extrusion by this mechanism was apparently not sustained. Following examination of all available geochronologic data from the eastern Himalayan, we suggest that the tectonics in this region alternated between one dominated by critical taper–frictional wedge tectonics and one dominated by channel flow–viscous wedge tectonics, over time spans of circa 2–5 Ma, from the Late Oligocene to the Late Miocene. Although Early Miocene tectonics are comparable across the Himalayan orogen, a fundamental change in the pattern of exhumation occurred in the Middle Miocene between the eastern Himalaya and more western regions in the orogen that is yet to be fully understood and incorporated into geodynamic models for orogenesis.

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