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Suppression of rice methane emission by sulfate deposition in simulated acid rain

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[1] Sulfate in acid rain is known to suppress methane (CH$_4$) emissions from natural freshwater wetlands. Here we examine the possibility that CH$_4$ emissions from rice agriculture may be similarly affected by acid rain, a major and increasing pollution problem in Asia. Our findings suggest that acid rain rates of SO$_4^{2-}$ deposition may help to reduce CH$_4$ emissions from rice agriculture. Emissions from rice plants treated with simulated acid rain at levels of SO$_4^{2-}$ consistent with the range of deposition in Asia were reduced by 24% during the grain filling and ripening stage of the rice season which accounts for 50% of the overall CH$_4$ that is normally emitted in a rice season. A single application of SO$_4^{2-}$ at a comparable level reduced CH$_4$ emission by 43%. We hypothesize that the reduction in CH$_4$ emission may be due to a combination of effects. The first mechanism is that the low rates of SO$_4^{2-}$ may be sufficient to boost yields of rice and, in so doing, may cause a reduction in root exudates to the rhizosphere, a key substrate source for methanogenesis. Decreasing a major substrate source for methanogens is also likely to intensify competition with sulfate-reducing microorganisms for whom prior SO$_4^{2-}$ limitation had been lifted by the simulated acid rain S deposition.


1. Introduction

[2] Methane (CH$_4$) is a potent greenhouse gas which is estimated to have contributed ~50% of the enhanced climate forcing of CO$_2$ since 1850 [Hansen et al., 2000]. Many of the sources of CH$_4$, both natural and anthropogenic, are reasonably well constrained within global budgets, the principal exception being rice agriculture emissions where estimates range from 20 to 100 Tg of the estimated total source of ~540 Tg [IPCC, 1996]. Despite tremendous growth in rice production due to ‘green revolution’ methods employed since the 1960s, recent estimates point to a rice CH$_4$ source at the lower end of this range [Denier van der Gon, 2000] with a projected likely declining trend in emissions as harvest area decreases and agricultural practices with respect to water management and fertilizer use change [Khalil and Shearer, 2006]. In rice agro-ecosystems such management practices are a key determinant of the output of CH$_4$. Methanogenesis is a strictly anaerobic process and so the duration of paddy inundation and the employment of mid-season field drainage both contribute to defining the maximum potential CH$_4$ that is produced versus that which is oxidised in the rhizosphere. This interplay between CH$_4$ production and oxidation may be further altered by the application of ammonium fertilizers [Bodelier et al., 2000] which may enhance the methanotrophic potential of paddy soil through reducing nitrogen limitation in methanotrophs [Bodelier et al., 2000].

[3] As well as nitrogen additions in the form of fertilizer, large individual applications of SO$_4^{2-}$ (10$^2$-10$^3$ kg SO$_4^{2-}$-S ha$^{-1}$) are increasingly being applied to rice paddies in order to meet the biological sulfate demand of rice plants. Such additions are known to suppress methane emissions from rice paddies by up to ~70% [Denier van der Gon et al., 2001, and references therein]. The application of large quantities of SO$_4^{2-}$ amendments to rice paddies has therefore been proposed as a greenhouse gas mitigation strategy [Denier van der Gon et al., 2001]. In natural peatlands, research has established that very low rates of SO$_4^{2-}$ deposition (~25 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ as small weekly pulses), similar to those of regions experiencing acid rain pollution, suppress methane emissions by as much as 30–40% [Dise and Verry, 2001; Gauci et al., 2002]. As with the effect of sulfate amendments on rice CH$_4$ emissions, it is thought that this is due to stimulation of sulfate-reducing microbial populations that out-compete methane producers for substrates [Gauci and Chapman, 2006]. For natural wetlands, this interaction may be important in offsetting the climate change-related growth in CH$_4$ emission from these ecosystems up until the middle of the 21st century [Gauci et al., 2004a].
[4] The effect of acid rain rates of \( \text{SO}_2^- \) introduced to wetlands may be long-lasting, with treatment effects continuing to suppress methane emissions (by 30–40\%) more than 2 years after the end of an acid rain simulation [Gauci et al., 2005]. This indicates that the available \( \text{SO}_2^- \) pool is continually being replenished through recycling at the steep redox gradients that exist primarily in the rhizosphere of these systems, such that more moles of \( \text{CH}_4 \) are suppressed by a mole of \( \text{SO}_2^- \) than the 1:1 reaction stoichiometry would suggest. Indeed up to 8.5 moles of \( \text{CH}_4 \) were suppressed per mole of applied acid rain simulant in this long-term Scottish peatland experiment [Gauci et al., 2005].

[5] Such recycling is also thought to occur in rice paddy soils where \( \text{SO}_2^- \) applications suppressed \( \text{CH}_4 \) emissions to a greater extent than the stoichiometric potential of the applied sulfate containing amendment would suggest [Denier van der Gon et al., 2001]. Given that rice \( \text{CH}_4 \) emissions seem to respond similarly to applied sulfate in terms of \( \text{CH}_4 \) emissions as do natural wetlands, the possibility arises that sulfate inadvertently applied through industrial pollutant emissions, atmospheric chemical transformation and subsequent deposition may similarly impact rice \( \text{CH}_4 \) emissions.

[6] Asia, as a consequence of rapid economic growth and limited pollution controls, has been a major and growing source of \( \text{SO}_2 \) to the atmosphere such that the region is currently contributing \( \sim 46\% \) of the global \( \text{SO}_2 \) source [Smith et al., 2001]. This is in contrast with higher latitude industrialised nations where \( \text{SO}_2 \) emissions have declined substantially since the 1970s. This global shift in anthropogenic sulfur emissions and deposition [Duan et al., 2007] means that Asian wetlands and rice agro-ecosystems will become the most impacted by S in the world and that this new source of S may significantly reduce \( \text{CH}_4 \) emissions from these ecosystems. Rice agriculture may be particularly affected since rice paddies are often not far removed from major industrial areas. Given that Asia produces the majority of the world’s rice and is experiencing this dramatic increase in acid rain pollution [Rodhe, 1999] we examine the potential effect of acid rain rates of \( \text{SO}_2^- \) deposition on \( \text{CH}_4 \) emission from rice agro-ecosystems.

[7] In this paper we detail the effect of an acid rain simulation experiment on rice mesocosm \( \text{CH}_4 \) emissions. In doing so we compare the effect of two modes and levels of sulfate application that are representative of sulfate applied either indirectly, through acid rain, or directly through the application of individual pulses of applied sulfate albeit within an annual acid rain S deposition range.

2. Methods

2.1. Rice Soils Collection and Mesocosm Preparation

[8] After considering \( \text{SO}_2^- \) deposition data we collected paddy soil prior to planting from Monte dos Alhos in Portugal (37.53°N, 8.27°W) in early 2004. In terms of S deposition, this area is relatively unpolluted (S deposition \( \sim 5 \) kg ha\(^{-1}\) yr\(^{-1}\)) since most deposition is from relatively clean air masses derived from the Atlantic Ocean.

[9] In order to approximate the \( \text{SO}_2^- \) concentration measured in the field, the following soil mix was prepared from the Portuguese soil: 50% coarse sand (horticultural grade, acid washed, leached with \( \text{dH}_2\text{O} \) to remove residual \( \text{SO}_2^- \)), 30% fine sand (horticultural grade sharp sand), 10% rice-paddy soil from Portugal, 5% peat and 5% part fine-ground silica. This diluted paddy soil ensured the presence of the same microbial communities that exist in rice soil and approximated the texture of the parent soil.

2.2. Rice Growing and Treatments

[10] Rice seeds from the Zeuz japonica cultivar (Oryza sativa L) were used for this experiment. This variety was chosen as it is commonly grown in the region from which the soil was collected. The rice seeds were first germinated in the dark on John Innes seed compost and then grown in trays until they reached a height of approximately 10 cm. The seedlings were then planted directly into rice soil mesocosms (4 plants per mesocosm 30 cm diameter \( \times 40 \) cm depth). The twelve mesocosms were fertilized with a balanced fertiliser (Miracle-Gro) at 8.48 g/mesocosm (applied diluted in \( \text{dH}_2\text{O} \)). This is equivalent to 1200 kg ha\(^{-1}\) of fertilizer and approximated N, P, K inputs annually applied to the original farm paddy soil. Rice plants were then grown under artificial lighting conditions (\( \sim 300 \) \( \mu \)E \( \text{m}^{-2} \) \( \text{h}^{-1} \)) that remained constant during the monitoring period i.e., early September to late November 2004. The mesocosms were randomly allocated to the following treatments (with 4 replicates per treatment): (1) regular small weekly applications of \( \text{Na}_2\text{SO}_4 \) at an annualised rate of 100 kg \( \text{SO}_2^- - \text{S} \) ha\(^{-1}\) yr\(^{-1}\), a rate of acid rain S deposition commonly experienced in NE China; (2) a single application of 100 kg \( \text{SO}_2^- - \text{S} \) ha\(^{-1}\) to replicate modes of sulfate application currently proposed as \( \text{CH}_4 \) mitigations strategies [Denier van der Gon et al., 2001]; and one set of controls where only deionised water and fertilizer were added. The small regular additions applied at a rate of 100 kg \( \text{SO}_2^- - \text{S} \) ha\(^{-1}\) yr\(^{-1}\) amounted to a total \( \text{SO}_2^- \) deposition of \( \sim 20 \) kg \( \text{SO}_2^- - \text{S} \) ha\(^{-1}\) throughout the duration of the study. A rhizon™ porous ceramic pore water sampler was positioned horizontally at 10 cm beneath the soil surface (Figure 1). This facilitated repeated (weekly) sampling of rhizosphere pore water for analysis of \( \text{SO}_2^- \) (via ion chromatography, Dionex) and for dissolved \( \text{CH}_4 \) [Gauci et al., 2004b].

2.3. \( \text{CH}_4 \) Flux Measurements

[11] Methane emissions were measured from each monolith once a week via a static chamber technique. In the middle part of each sampling day samples were taken from a headspace volume that was defined by temporarily sealing a chamber onto the monolith for 40 min (Figure 1). The chamber contained a fan which allowed good mixing of headspace gases during incubation. During the period of incubation, a 60 ml air sample was withdrawn from each chamber at 0, 20 and 40 min. Samples were analyzed for \( \text{CH}_4 \) content by GC FID and \( \text{CH}_4 \) fluxes were calculated. A chamber extension collar was used to minimize damage to plants once they had reached the maximum height of the chamber.

[12] Ancillary measurements made included soil and air temperature; plant height; tiller number; root, shoot, and grain dry-weight; total C, N, S of root, shoot, and grain and plant morphological observations (e.g., timing of spikelet formation, filling of panicles etc). Results were analyzed for
statistical differences using repeated measures analysis of variance (General MANOVA; Minitab release 14).

3. Results and Discussion

Over the duration of the experiment, CH$_4$ emission was lower from both sulfate treatments relative to controls. This decrease was significant for the single application of 100 kg S ha$^{-1}$ treatment, where CH$_4$ emission was reduced by up to one third ($p < 0.01$; Table 1). Methane emissions from control mesocosms followed a pattern of almost linear increase beyond day 14, progressively increasing in CH$_4$ output until reaching a peak in emission in day 56 of 174 mg CH$_4$ m$^{-2}$ day$^{-1}$ (Figure 2). In contrast, both of the sulfate treatments demonstrated reductions in CH$_4$ emission growth between days 28 and 38, and lower CH$_4$ emissions during the peak emission period exhibited by the controls (121 mg CH$_4$ m$^{-2}$ day$^{-1}$ in the acid rain simulation and 99 mg CH$_4$ m$^{-2}$ day$^{-1}$ in the single sulfate pulse mesocosms) (Figure 2). Broad patterns in CH$_4$ emission growth reflect changes in pore water dissolved CH$_4$ and SO$_4^{2-}$ concentration (Figure 3). For the large single SO$_4^{2-}$ application treatment the onset of growth in CH$_4$ production is delayed relative to controls and the acid rain simulation treatment until after day 35 (Figure 3b). Dissolved CH$_4$ concentrations peaked at 120 μmol L$^{-1}$ for both controls and acid rain simulation treatment on day 70 but only peaked at 80 μmol L$^{-1}$ in the large single SO$_4^{2-}$ application. Dissolved CH$_4$ concentrations were significantly lower in the large single SO$_4^{2-}$ application than in either the controls or the acid rain simulation ($p < 0.001$, General MANOVA) (Figure 3b).

Dissolved SO$_4^{2-}$ concentrations were an order of magnitude higher in the single SO$_4^{2-}$ application mesocosms than in the control and acid rain simulation mesocosms during the first 30 d. Concentrations peaked on day 14, reflecting the period required for the dissolved SO$_4^{2-}$ to percolate down to the rhizone pore water collection point 10 cm below the soil sediment/paddy water interface (Figure 3a). As with dissolved CH$_4$ concentrations there was no difference between control and acid rain SO$_4^{2-}$ simulation treatments throughout the duration of the monitoring period. This may not reflect differences in situ sulfate reduction rates since increased sulfate-reducing activity may accelerate S cycling such that instantaneous SO$_4^{2-}$ concentrations are unaffected or may even be reduced [Gauci et al., 2002; Gauci and Chapman, 2006]. Given that pore water concentrations were measured at only one depth, the possibility remains that any differences in CH$_4$ concentration between the acid rain simulation treatments and controls were located elsewhere in the mesocosm soil profile.

Table 1. Summary Table of CH$_4$ Emission Response to Experimental SO$_4^{2-}$ Addition$^a$

<table>
<thead>
<tr>
<th>Treatment</th>
<th>Mean CH$_4$ Flux (±SE) (mg CH$_4$ m$^{-2}$ d$^{-1}$)</th>
<th>Total CH$_4$ Emission (g CH$_4$ m$^{-2}$)</th>
<th>Percent Suppression (Treatment Versus Control)</th>
<th>P-value (MANOVA)</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Duration of Experiment (Day 1 – 77) n = 60</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Control</td>
<td>82.9 (8.8)</td>
<td>6.61</td>
<td></td>
<td>n.s.</td>
</tr>
<tr>
<td>100 kg S ha$^{-1}$</td>
<td>75.3 (6.8)</td>
<td>5.74</td>
<td>13</td>
<td>&lt;0.01</td>
</tr>
<tr>
<td>100 kg S ha$^{-1}$</td>
<td>64.0 (8.8)</td>
<td>4.86</td>
<td>31</td>
<td>&lt;0.01</td>
</tr>
<tr>
<td><strong>Grain Filling/Ripening Stage (Day 38 – 63) n = 20</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Control</td>
<td>134 (14.8)</td>
<td>3.65</td>
<td></td>
<td>&lt;0.05</td>
</tr>
<tr>
<td>100 kg S ha$^{-1}$</td>
<td>103 (8.8)</td>
<td>2.77</td>
<td>24</td>
<td>&lt;0.05</td>
</tr>
<tr>
<td>100 kg S ha$^{-1}$</td>
<td>86 (10.5)</td>
<td>2.45</td>
<td>43</td>
<td>&lt;0.05</td>
</tr>
</tbody>
</table>

$^a$Here $P$ values denote the significance level of any difference between individual treatments and the control.
The observed period of reduced CH$_4$ emission growth rate in SO$_4^{2-}$/C$_0$ treated mesocosms coincides with the first appearance of panicles on the rice plants (day 28) and continued until all panicles had formed and flowered (by day 38). During the ensuing grain-filling stage, the cumulative small applications of sulfate in the acid rain simulation were sufficient to suppress emissions by 24% (p < 0.05; Table 1), an important reduction in CH$_4$ emission given that this period constituted ~50% of the total CH$_4$ that was emitted during the experiment. During this period, CH$_4$ emission from the single 100 kg S ha$^{-1}$ treatment was significantly suppressed by 43% relative to controls (p < 0.05; Table 1). Suppressed emitted CH$_4$:SO$_4^{2-}$ ratios were, over the entire experiment, 26 and 11 for the continuous and single sulfate application modes respectively, and 32 and 7.5 for the same treatments during grain filling and ripening. For the small continuous SO$_4^{2-}$ applications these are the highest ratios recorded for terrestrial soil CH$_4$ emission studies, implying extensive recycling of S in the rhizosphere should all CH$_4$ suppression be attributable to this single mechanism.

The timing of the maximum period of suppression (grain filling and ripening) suggests that substrate supply via rhizodeposition of photosynthetically fixed carbon may be a major control in determining the extent of competition between sulfate-reducing bacteria and methanogens. Given that the hypothesized interaction between sulfate-reducing bacteria and methanogens is dependent on a limiting supply of substrate over which competitive exclusion of methanogens may occur, this period of maximum suppression may be taking place during a period of reduced rhizodeposition. Grain filling and ripening requires a substantial proportion

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**Figure 2.** Methane emission from SO$_4^{2-}$ treated and control rice mesocosms. Each point represents the mean CH$_4$ flux from four mesocosms. Error bars represent ±1 SE of the mean.

**Figure 3.** (a) Dissolved SO$_4^{2-}$ concentrations in pore water from the two modes of SO$_4^{2-}$ treatment and controls. Points represent the mean of four measurements. Error bars are omitted for clarity. (b) Dissolved CH$_4$ concentration in mesocosm pore water (sampled at 10 cm beneath the soil surface/paddy water interface) for the two modes of SO$_4^{2-}$ treatment and controls.
impacted regions such as Asia is capable of substantially suppressing CH$_4$ emissions. Significant CH$_4$ suppression was measured during the rice grain filling and ripening stage of plant development, which constitutes $\sim$50% of the total CH$_4$ that is emitted by controls (Table 1). This temporal feature, together with a trend of increasing harvest index with increasing applied total SO$_4^{2-}$, suggests that SO$_4^{2-}$ influences over rhizodeposited exudate supply may exert an overall control on methanogens via 2 mechanisms: (1) the limitation of substrate supply while (2) simultaneously providing a more competitive environment in which competition with sulfate reducers intensifies.

[19] Estimated mean total annual SO$_4^{2-}$ deposition in the region ranges from $\sim$30 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ in S. East Asia to $\sim$85 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ in East Asia [Dentener et al., 2006] although areas within East Asia (e.g., the North) may receive considerably more S deposition. These values are similar to the total SO$_4^{2-}$ applied in both the small continuous acid rain simulation ($\sim$20 kg SO$_4^{2-}$-S ha$^{-1}$) and in the single application of 100 kg SO$_4^{2-}$-S ha$^{-1}$. Our experimental approach denied SO$_4^{2-}$ treated mesocosms a history of aerially deposited SO$_4^{2-}$ and so represents the situation at the outset of a pollution episode. Given that periods of drought in natural wetlands (an analogous situation to regular dry periods between periods of inundation in rice agriculture) are known to replenish the supply of SO$_4^{2-}$ from reduced S forms [Dowrick et al., 2006], the CH$_4$ response to acid rain S deposition may, in nature, respond to cumulative S deposition over many years even if a component of the deposited SO$_4^{2-}$ was to be dispersed in drainage waters. Our findings may therefore underestimate the overall CH$_4$ suppression response to prolonged periods of acid deposition and SO$_4^{2-}$ replenishment through periodic paddy drainage.

[20] By 2030 pollutant S deposition is predicted to significantly worsen with deposition under the most pessimistic IPCC Special Report on Emissions Scenarios (SRES) scenario ranging from $\sim$88 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ in S. East Asia to $\sim$130 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ in East Asia and, highest, $\sim$150 kg SO$_4^{2-}$-S ha$^{-1}$ yr$^{-1}$ in South Asia [Dentener et al., 2006], well within the range of S application shown to reduce CH$_4$ emissions [Denier van der Gon et al., 2001] and therefore negating the need for costly SO$_4^{2-}$ amendments to meet the biological demands of rice plant growth. A repeat of this study for a range of commonly grown rice cultivars and commonly farmed soil types under different management regimes would ensure that the full response range of rice CH$_4$ emission to acid rain is properly characterized for Asia.

[21] In conclusion, our results point to a hitherto overlooked factor that, in addition to recognized changes in management practices [Khalil and Shearer, 2006], may be progressively reducing CH$_4$ emissions from rice agriculture. Our findings suggest that acid rain rates of SO$_4^{2-}$ deposition may help to reduce CH$_4$ emissions from rice through a combination of mechanisms. The first mechanism is that the low rates of SO$_4^{2-}$ supply may be sufficient to boost yields of rice and, in so doing, causes a reduction in root exudates to the rhizosphere, a key substrate source for methanogenesis. The role of substrate limitation is further supported by the observation that SO$_4^{2-}$ suppression of CH$_4$ emission was greatest during the energy demanding grain production stage of plant development, which constitutes $\sim$50% of the total CH$_4$ that is emitted by controls (Table 1). This temporal feature, together with a trend of increasing harvest index with increasing applied total SO$_4^{2-}$, suggests that SO$_4^{2-}$ influences over rhizodeposited exudate supply may exert an overall control on methanogens via 2 mechanisms: (1) the limitation of substrate supply while (2) simultaneously providing a more competitive environment in which competition with sulfate reducers intensifies.

Figure 4. Changes in rice harvest index (HI) with applied SO$_4^{2-}$. Error bars represent $\pm$1 SE of the mean. Bars with the same letter do not differ at the p < 0.05 level, ANOVA.

4. Summary and Implications

[18] Our experimental findings demonstrate that SO$_4^{2-}$ applied at rates within the range experienced in acid rain
and ripening stage. This suggests that photosynthesis leakage to the rhizosphere was reduced owing to the greater energy demands required by reproduction. Decreasing a major substrate source for methanogens is likely to fuel competition with sulfate-reducing microorganisms for whom $\text{SO}_4^{2-}$ availability had been limiting prior to the onset of simulated acid rain $\text{S}$ deposition.

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References

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