Organic geochemistry of the crater-fill sediments from Boltysh impact crater, Ukraine

Conference Item

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INTRODUCTION

The Boltysh crater has been known for several decades and was first drilled in the 1960s as part of a study of economic oil shale deposits. Unfortunately, the cores were not curated and have been lost. We have re-drilled the impact crater and have recovered a near continuous record of ~400 m of organic-rich sediments together with 15 m of suevite. The sediments were deposited in a deep isolated lake and span a period ~10 Ma.

The Boltysh impact crater, centred at 48°54'N and 32°15'E, is a complex impact structure formed on the crystalline basement rocks of the Ukrainian shield, which comprise porphyroblastic granites (age ca. 1.55 Ga) and biotite gneisses (age ca. 1.85 – 2.22 Ga) [1]. The structure is covered by Quaternary sediments and has been dated at 65.17±0.64 Ma [2]. At 24km diameter, the impact is unlikely to have contributed substantially to the worldwide devastation at the end of the Cretaceous.

PRELIMINARY RESULTS

The lowermost 15m of the core is predominantly a polymict suevite breccia with evidence of secondary hydrothermal activity (Figure 2a). A sharp, 60°, angular contact marks the boundary between the impact breccia and the overlying sedimentary deposits (Figure 2b). The first sediments to be deposited in the crater lake occur at 581.5m and comprise a series of thin turbidite beds overlain by progressively more organic-rich shales (Figure 3). The 400 m of overlying sediments cover a period of approximately 10Ma. Preliminary palynological investigation indicates a number of significant floral and faunal transitions throughout the core and work is in progress to establish if the Cretaceous-Tertiary Boundary exists within the basal section of the core.

Distribution of organic compounds indicate the main sources of organic matter in the post-impact sediments are dominated by algal input (Figure 4) with an increasing higher plant contribution up the core, this trend is also observed in the C/N ratios. From isomeric distributions of biomarkers, sediments just above the impact boundary are still thermally immature (before the oil generation window). However there is evidence for some heating/warming of the overlying of the sediments from the residual heat from the impact, but only in 12 m of sediment above the impact boundary (Figure 5). Carbon isotopic analysis of the bulk organic material shows that there are a number of pronounced isotopic excursions in the first 100 m of core above the boundary (Figure 6).

PLANNED WORK

Work is in progress to complete the detailed palynological survey of the core and organic geochemical analyses. These studies will enable us to reconstruct the paleoenvironmental history of the post-impact environment at Boltysh and to examine the subsequent paleoenvironmental record preserved in the crater-fill sediments for a region of North Tethys where paleoenvironmental information is presently scarce.

REFERENCES


Figure 1: Location of Boltysh impact crater and estimated range of ejector layer

Figure 2: A) Suevite containing glass veins and B) Boundary between impact breccia (on left) and overlying sediments.

Figure 3: Sedimentary log of bottom 100 m of Boltysh core.

Figure 4: Partial mass chromatogram of sample at 568.6 m highlighting algal input

Figure 5: Ratio of $bb/(ab + ba + bb)$ C31 hopanes indicating limited heating of sediments post impact

Figure 6: Bulk organic $\delta^{13}$C values for crater fill sediments